



Glacial Cycles

Richard McGehee



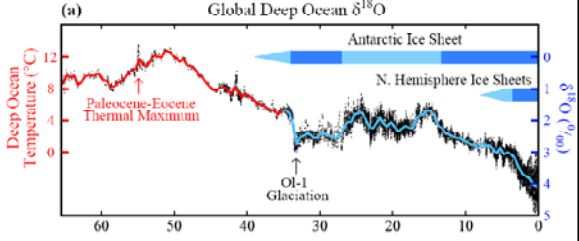
Seminar on the Mathematics of Climate Change
School of Mathematics
October 7, 2009

Glacial Cycles




Temperatures in the Cenozoic Era

Global Deep Ocean $\delta^{18}\text{O}$

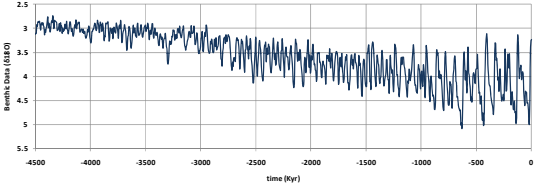


Hansen, et al, Target atmospheric CO₂: Where should humanity aim? *Open Atmos. Sci. J.* 2 (2008)

Glacial Cycles




^{18}O in Foraminifera Fossils During the Past 4.5 Myr

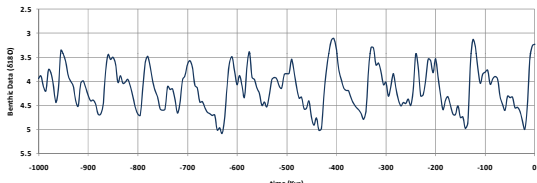


Lisiecki, L. E., and M. E. Raymo (2005), A Pliocene-Pleistocene stack of 57 globally distributed benthic $\delta^{18}\text{O}$ records, *Paleoceanography*, 20, PA1003, doi:10.1029/2004PA001071.

Glacial Cycles




^{18}O in Foraminifera Fossils During the Past 1.0 Myr



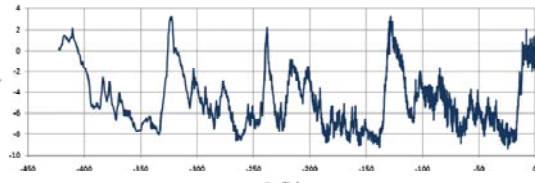
Lisiecki, L. E., and M. E. Raymo (2005), A Pliocene-Pleistocene stack of 57 globally distributed benthic $\delta^{18}\text{O}$ records, *Paleoceanography*, 20, PA1003, doi:10.1029/2004PA001071.

Glacial Cycles




Recent (last 400 Kyr) Temperature Cycles

Vostok Ice Core Data



J.R. Petit, et al (1999) Climate and atmospheric history of the past 420,000 years from the Vostok ice core, Antarctica, *Nature* 399, 429-436.

Glacial Cycles



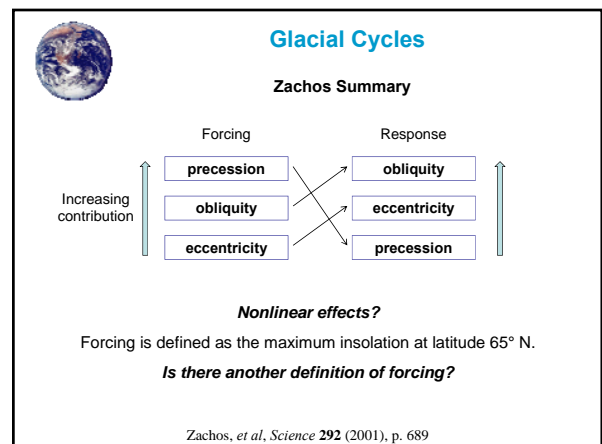
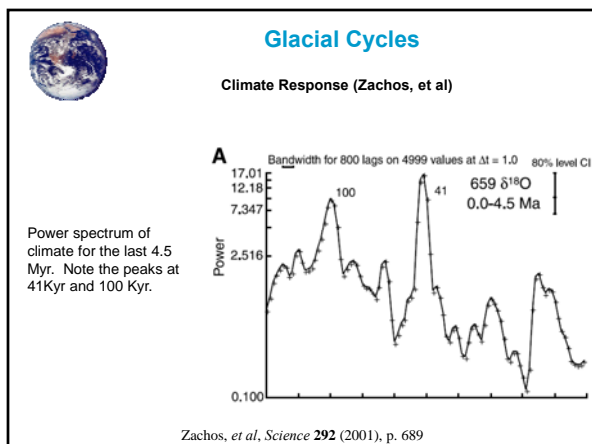
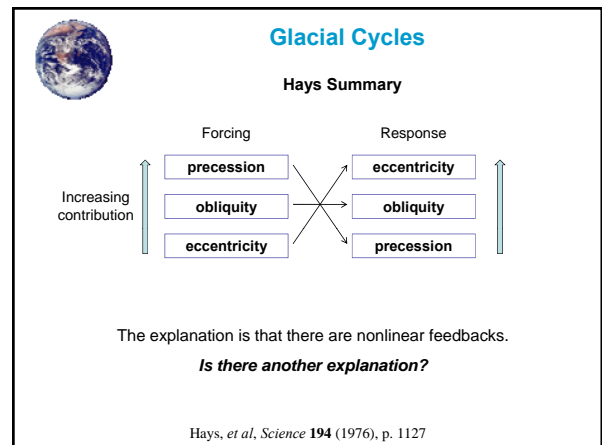
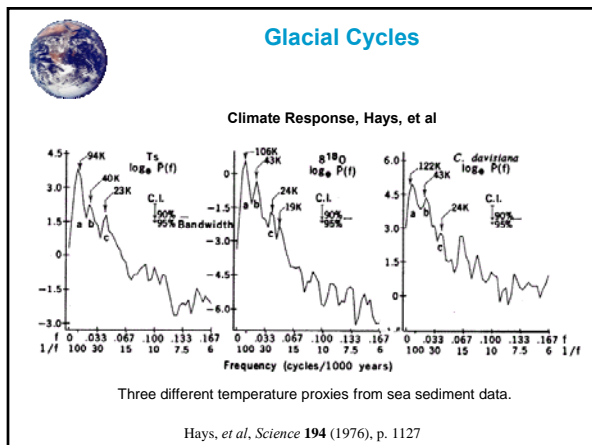
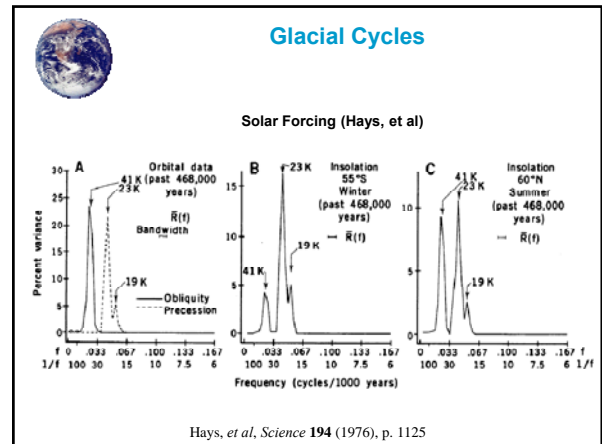
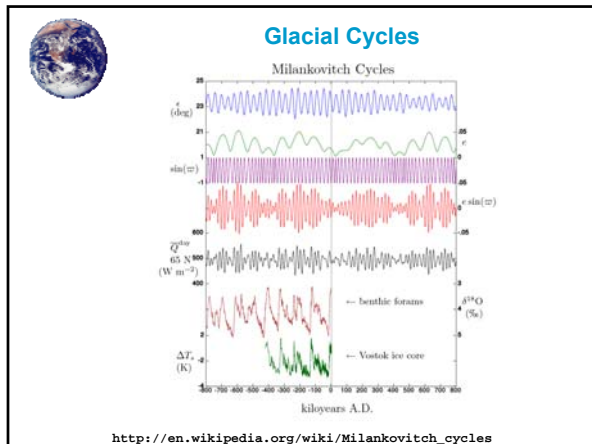
What Causes Glacial Cycles?

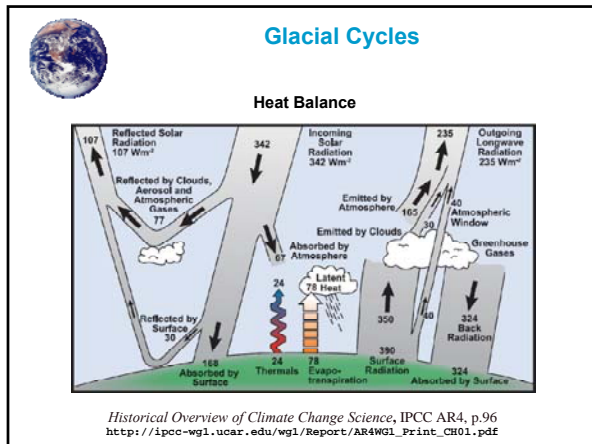
Widely Accepted Hypothesis

The glacial cycles are driven by the variations in the Earth's orbit (Milankovitch Cycles), causing a variation in incoming solar radiation (insolation).

This hypothesis is widely accepted, but also widely regarded as insufficient to explain the observations.

The additional hypothesis is that there are feedback mechanisms that amplify the Milankovitch cycles. What these feedbacks are and how they work is not fully understood.





Glacial Cycles

Ice Albedo Feedback Model

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

$Qs(y)$ = insolation $(1 - \alpha(y, \eta))$ = albedo $(A + BT)$ = outward radiation $C(\bar{T} - T)$ = heat transport

y = sine of latitude
 $T(y)$ = annual mean temperature
 $Qs(y)$ = annual mean insolation
 Q = global annual mean insolation

This equation has a stable equilibrium consisting of polar ice caps.
 The latitude of the equilibrium ice boundary and the equilibrium global annual mean temperature are functions of the parameters.

K.K. Tung, *Topics in Mathematical Modeling*, Princeton (2007), Chapt 8

Glacial Cycles

Ice Albedo Feedback Model

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

Idea

Instead of solar forcing (maximum insolation at 65 N latitude), use the global annual mean temperature predicted by the model.

Using Kepler's Laws, we can compute:

$$Q = \frac{Ka^2}{\sqrt{1-e^2}}$$

$$s(y) = \frac{2}{\pi^2} \int_0^{2\pi} \sqrt{1 - (\sqrt{1-y^2} \sin \beta \cos \gamma - y \cos \beta)^2} d\gamma$$

a = semimajor axis
 e = eccentricity
 β = obliquity

Glacial Cycles

Ice Albedo Feedback Model

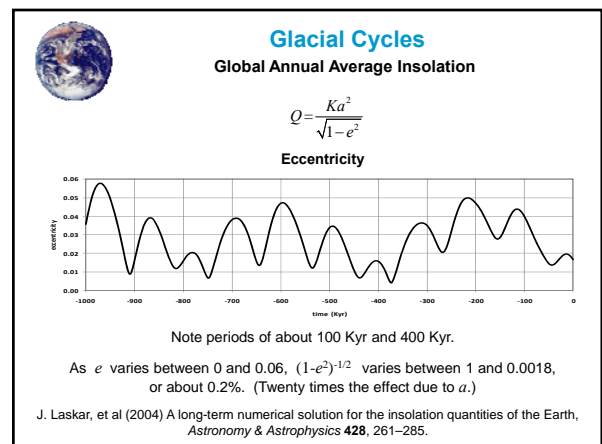
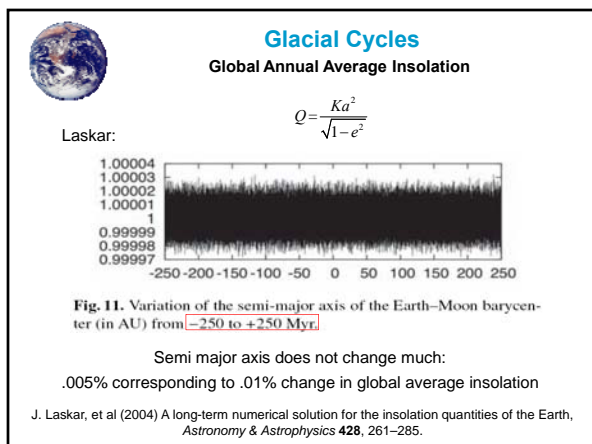
$$Q = \frac{Ka^2}{\sqrt{1-e^2}}$$

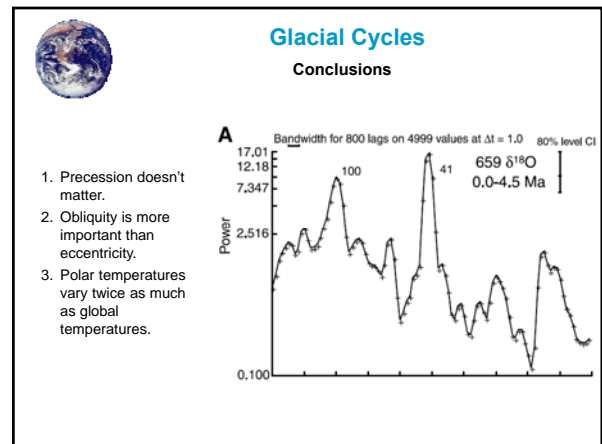
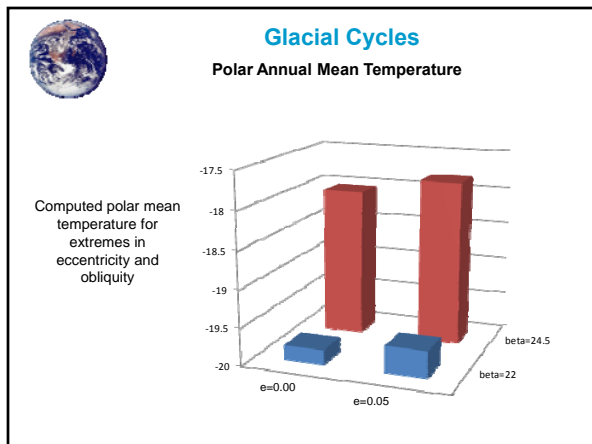
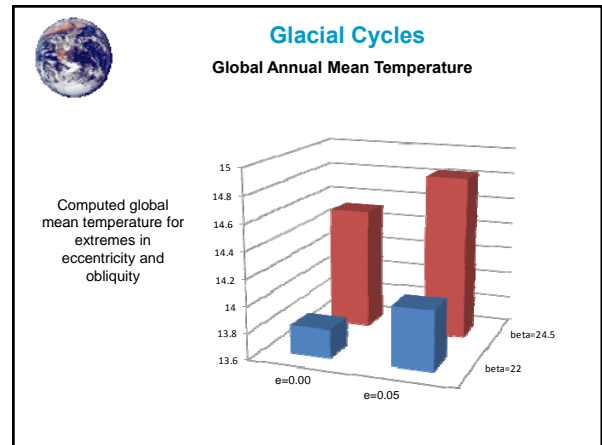
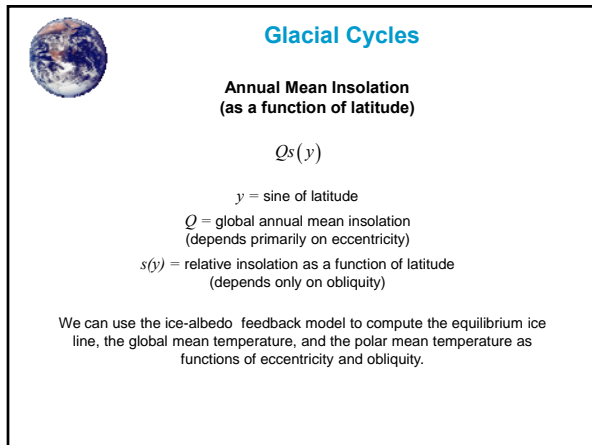
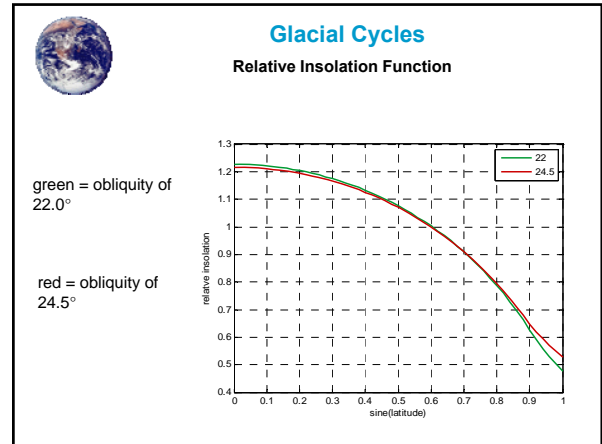
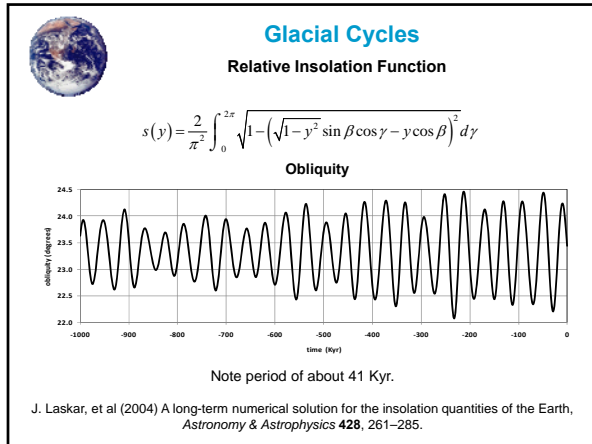
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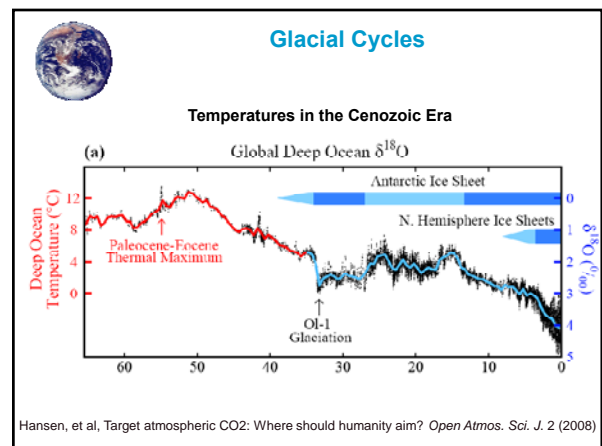
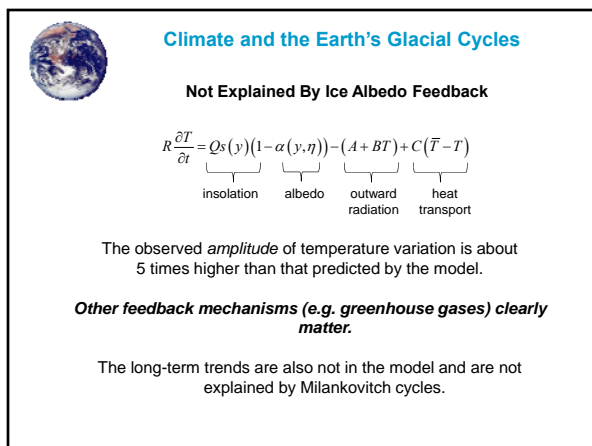
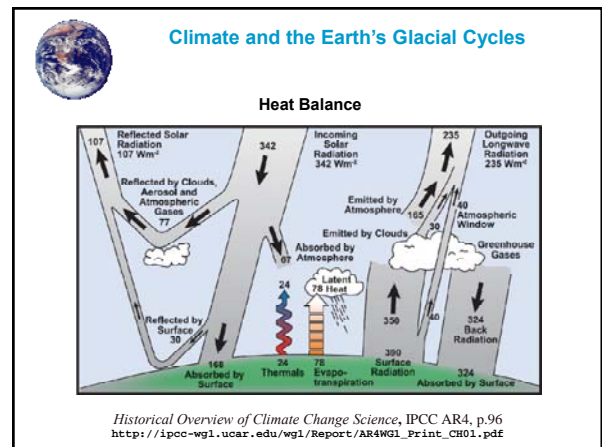
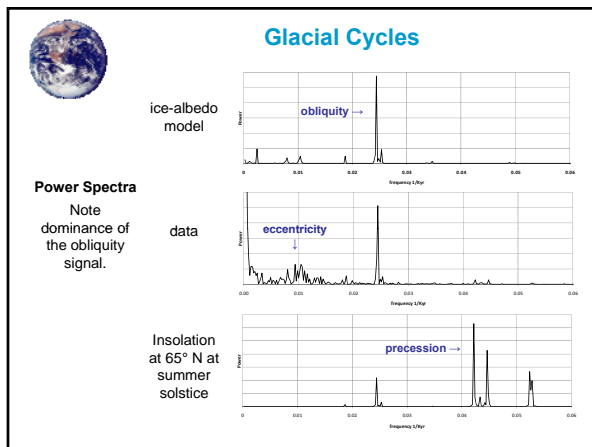
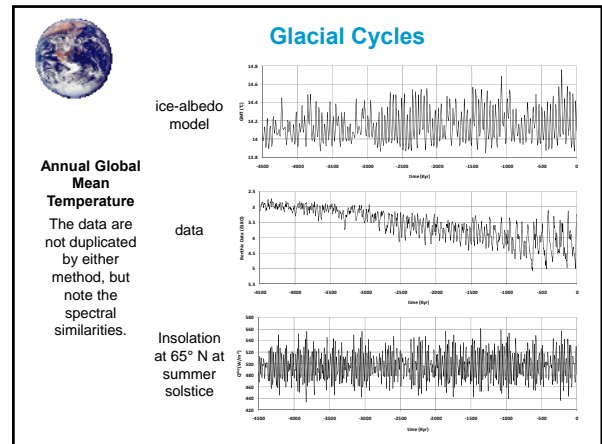
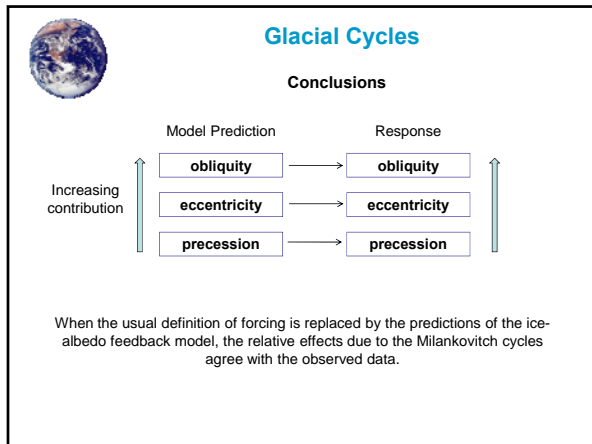
Note that Q , the global annual mean insolation depends only on the semimajor axis and the eccentricity.

Note that $s(y)$, the insolation distribution by latitude, depends only on the obliquity.

Note that the effect due to precession disappears when averaged over a year.





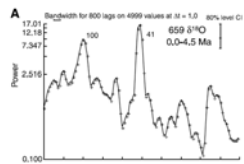




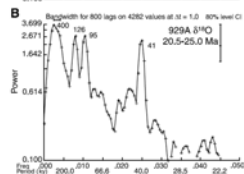
Glacial Cycles

More from Zachos

A. Power spectrum of climate for the last 4.5 Myr. Note the peaks at 41Kyr and 100 Kyr.



B. Power spectrum of climate for the period 25 Myr bp to 20.5 Myr bp. Note the new peak at 400 Kyr and the "split" peaks at 126Kyr and 95 Kyr.



Zachos, *et al*, *Science* **292** (2001), p. 689