



A Finite Dimensional Model of Ice-Albedo Feedback

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


Ice-Albedo Feedback

Budyko-Sellers Model

$$R \frac{\partial T}{\partial t} = \underbrace{Qs(y)}_{\text{insolation}} (1 - \underbrace{\alpha(y, \eta)}_{\text{albedo}}) - \underbrace{(A + BT)}_{\text{re-radiation}} + C(\underbrace{\bar{T} - T}_{\text{transport}})$$

$T = T(y, t)$: annual mean surface temperature
 $y = \sin(\text{latitude}) \quad y \in [0, 1]$
 Q : global annual mean insolation
 $s(t)$: relative annual mean insolation $\int_0^1 s(y) dy = 1$
 $y = \eta$: ice boundary
 $\alpha(y, \eta) = \begin{cases} \alpha_1, & y < \eta, \\ \alpha_2, & y > \eta. \end{cases}$ albedo
 $\bar{T}(t) = \int_0^1 T(y, t) dy$: global annual mean temperature



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$Q = 343 \text{ W/m}^2 \quad A = 202 \text{ W/m}^2$
 $B = 1.9 \text{ W/m}^2/\text{C} \quad C = 3.04 \text{ W/m}^2/\text{C}$
 $\alpha_1 = 0.32$: (water and land)
 $\alpha_2 = 0.62$: (ice)

y = proportion of Earth's surface between latitude $-\arcsin(y)$ and $+\arcsin(y)$,
 η = proportion of Earth's surface that is ice-free.

If σ = surface area of Earth ($\approx 5.1 \times 10^{14} \text{ m}^2$), then
 $\sigma(1-\eta)$ = surface area of ice in square meters.




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Heat Capacity

R = heat capacity of Earth's surface ($\text{J/m}^2/\text{C}$).
 Heat capacity of water $\approx 4 \text{ J/g}^\circ\text{C}$.
 Mass of 1 cm^3 of $\text{H}_2\text{O} \approx 1 \text{ g}$.
 Assumption: Earth's surface consists of 100 m of H_2O .
 Each square meter of Earth's surface consists of
 $100 \text{ m}^3 = 10^8 \text{ cm}^3$ of H_2O , or about 10^8 g , so
 $R \approx 4 \times 10^8 \text{ J/m}^2/\text{C}$.




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$$R \frac{\partial T}{\partial t} = \underbrace{Qs(y)}_{\text{insolation}} (1 - \underbrace{\alpha(y, \eta)}_{\text{albedo}}) - \underbrace{(A + BT)}_{\text{re-radiation}} + C(\underbrace{\bar{T} - T}_{\text{transport}})$$

Time Scale

A Watt is a Joule per **second**, but all quantities are **annual** averages.
 One year is approximately $\kappa = 3.16 \times 10^7$ seconds.
 If we measure time in years instead of seconds, the equation becomes

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T))$$


Ice-Albedo Feedback

Budyko-Sellers-Widiasih Model

$$\frac{d\eta}{dt} = \epsilon \kappa (T_s - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T))$$


$$T_s = \frac{T(\eta^-) + T(\eta^+)}{2}$$

$$\bar{T} = \bar{T}(t) = \int_0^1 T(y, t) dy$$

What about ϵ ?

The parameter ϵ should be determined from climate data. (good project!)

Best current guess:
 $\epsilon \approx 10^{-12}$

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Budyko-Sellers-Widiasih Model

Heat of Fusion

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

To move the ice line, we must melt or freeze water.
Heat of fusion of water: 334 J/g, or 3.34×10^8 J/m³


Assumption: Average thickness of ice is 450 m.
Energy to melt 1 m² of ice: $\Omega = 1.5 \times 10^{11}$ Joules
Energy to move ice line from η to $\eta + \Delta\eta$:

$$\Omega\sigma\Delta\eta \text{ Joules, or } \Omega\Delta\eta \text{ J/m}^2$$

(σ = surface area of Earth.)

Energy needed to move the ice line:

$$\Omega \frac{d\eta}{dt} \text{ J/yr/m}^2$$

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Budyko-Sellers-Widiasih Model


$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T) - \varepsilon\Omega(T_b - T_c))$$

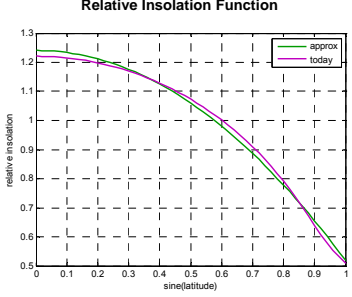
State Space

η lives in $[0, 1]$.
 T lives in a space of functions on $[0, 1]$.

What space of functions?


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Relative Insolation Function



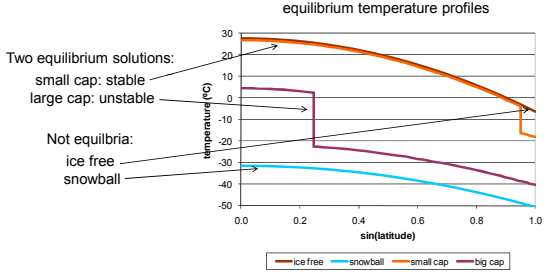
green = quadratic approximation (Tung and North)

mauve = formula using obliquity of 23.5°

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If s is assumed to be quadratic, then the equilibrium solutions are piecewise quadratic with a discontinuity at the ice boundary.

equilibrium temperature profiles



Two equilibrium solutions:
small cap: stable
large cap: unstable

Not equilibria:
ice free
snowball

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Budyko-Sellers-Widiasih Model

$$\frac{d\eta}{dt} = \varepsilon\kappa(T_b - T_c)$$

$$\frac{\partial T}{\partial t} = \frac{\kappa}{R} (Q_s(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T) - \varepsilon\Omega(T_b - T_c))$$


State Space

What space of functions?

Assume that T lives in the space of functions on $[0, 1]$ which are piecewise quadratic with a discontinuity at the ice boundary.

$$T(y) = \begin{cases} U(y), & y < \eta, \\ V(y), & y > \eta, \\ (U(\eta) + V(\eta))/2, & y = \eta, \end{cases}$$

where U and V are quadratic on $[0, 1]$.

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Budyko-Sellers-Widiasih Model

Legendre Polynomials

Since s is even, we assume that U and V are even and expand using the first two even Legendre polynomials:

$$p_0(y) = 1$$

$$p_2(y) = \frac{1}{2}(3y^2 - 1)$$

$$U(y, t) = u_0(t)p_0(y) + u_2(t)p_2(y) = u_0(t) + u_2(t)p_2(y),$$

$$V(y, t) = v_0(t)p_0(y) + v_2(t)p_2(y) = v_0(t) + v_2(t)p_2(y),$$

$$s(y) = s_0p_0(y) + s_2p_2(y) = 1 + s_2p_2(y).$$

$$\frac{\partial U}{\partial t} = \dot{u}_0 + \dot{u}_2 p_2(y), \quad \left(\dot{u} = \frac{du}{dt} \right)$$

$$\frac{\partial V}{\partial t} = \dot{v}_0 + \dot{v}_2 p_2(y).$$



Ice-Albedo Feedback

Five Dimensional Budyko-Sellers-Widiasih Model

Summary

$$\begin{aligned} \dot{\eta} &= \varepsilon \kappa (T_b - T_c) \\ \dot{u}_0 &= \frac{\kappa}{R} (Q(1 - \alpha_1) - A + C\bar{T} - (B + C)u_0 - \varepsilon \Omega (T_b - T_c)) \\ \dot{v}_0 &= \frac{\kappa}{R} (Q(1 - \alpha_2) - A + C\bar{T} - (B + C)v_0 - \varepsilon \Omega (T_b - T_c)) \\ \dot{u}_2 &= \frac{\kappa}{R} (Qs_2(1 - \alpha_1) - (B + C)u_2) \\ \dot{v}_2 &= \frac{\kappa}{R} (Qs_2(1 - \alpha_2) - (B + C)v_2) \\ T_b &= \frac{u_0 + v_0}{2} + \frac{u_2 + v_2}{2} p_2(\eta) \\ \bar{T} &= \eta u_0 + (1 - \eta)v_0 + p_2(\eta)(u_2 - v_2) \end{aligned}$$



Ice-Albedo Feedback

Five Dimensional Budyko-Sellers-Widiasih Model

Quadratic Modes

$$\begin{aligned} \dot{u}_2 &= \frac{\kappa}{R} (Qs_2(1 - \alpha_1) - (B + C)u_2) \\ \dot{v}_2 &= \frac{\kappa}{R} (Qs_2(1 - \alpha_2) - (B + C)v_2) \end{aligned}$$

The dynamics of u_2 and v_2 are independent of each other and of the other variables (including η). These variables decay exponentially to

$$u_2^* = \frac{Qs_2(1 - \alpha_1)}{(B + C)} \quad v_2^* = \frac{Qs_2(1 - \alpha_2)}{(B + C)}$$

Exercise: All higher order modes behave the same way.



Ice-Albedo Feedback

Three Dimensional Budyko-Sellers-Widiasih Model

Globally Attracting 3D Subspace Summary

$$\begin{aligned} \dot{\eta} &= \varepsilon \kappa (T_b - T_c) \\ \dot{u}_0 &= \frac{\kappa}{R} (Q(1 - \alpha_1) - A + C\bar{T} - (B + C)u_0 - \varepsilon \Omega (T_b - T_c)) \\ \dot{v}_0 &= \frac{\kappa}{R} (Q(1 - \alpha_2) - A + C\bar{T} - (B + C)v_0 - \varepsilon \Omega (T_b - T_c)) \end{aligned}$$

Ice Boundary Temperature

$$T_b = \frac{u_0 + v_0}{2} + p_2(\eta) \frac{Qs_2(1 - \alpha_0)}{B + C}$$

Global Mean Temperature

$$\bar{T} = \eta u_0 + (1 - \eta)v_0 + p_2(\eta) \frac{Qs_2(\alpha_2 - \alpha_1)}{B + C}$$



Ice-Albedo Feedback

Three Dimensional Budyko-Sellers-Widiasih Model

$$\begin{aligned} \dot{u}_0 &= \frac{\kappa}{R} (Q(1 - \alpha_1) - A + C\bar{T} - (B + C)u_0 - \varepsilon \Omega (T_b - T_c)) \\ \dot{v}_0 &= \frac{\kappa}{R} (Q(1 - \alpha_2) - A + C\bar{T} - (B + C)v_0 - \varepsilon \Omega (T_b - T_c)) \end{aligned}$$

Further Reduction

$$w = \frac{u_0 + v_0}{2}, \quad z = u_0 - v_0$$

$$\dot{w} = \frac{\dot{u}_0 + \dot{v}_0}{2} = \frac{\kappa}{R} (Q(1 - \alpha_0) - A + C\bar{T} - (B + C)w - \varepsilon \Omega (T_b - T_c))$$

$$\dot{z} = \dot{u}_0 - \dot{v}_0 = \frac{\kappa}{R} (Q(\alpha_2 - \alpha_1) - (B + C)z)$$



Ice-Albedo Feedback

Three Dimensional Budyko-Sellers-Widiasih Model

$$\begin{aligned} \dot{\eta} &= \varepsilon \kappa (T_b - T_c) \\ \dot{w} &= \frac{\kappa}{R} (Q(1 - \alpha_0) - A + C\bar{T} - (B + C)w - \varepsilon \Omega (T_b - T_c)) \\ \dot{z} &= \frac{\kappa}{R} (Q(\alpha_2 - \alpha_1) - (B + C)z) \end{aligned}$$

Note that z converges exponentially to its equilibrium value

$$z^* = \frac{Q(\alpha_2 - \alpha_1)}{B + C}$$

Attracting invariant 2D subspace

$$\begin{aligned} \dot{\eta} &= \varepsilon \kappa (T_b - T_c) \\ \dot{w} &= \frac{\kappa}{R} (Q(1 - \alpha_0) - A + C\bar{T} - (B + C)w - \varepsilon \Omega (T_b - T_c)) \end{aligned}$$



Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

Ice Boundary Temperature


$$T_b = \frac{u_0 + v_0}{2} + p_2(\eta) \frac{Qs_2(1 - \alpha_0)}{B + C}$$

Recall

$$w = \frac{u_0 + v_0}{2}, \quad z = u_0 - v_0$$

$$T_b = T_b(\eta, w) = w + p_2(\eta) \frac{Qs_2(1 - \alpha_0)}{B + C}$$

$$T_b(\eta, w) = w + \frac{Q(1 - \alpha_0)}{B + C} s_2 p_2(\eta)$$



Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

Global Mean Temperature

$$\bar{T} = \eta u_0 + (1-\eta)v_0 + P_2(\eta) \frac{Qs_2(\alpha_2 - \alpha_1)}{B+C}$$


Recall

$$w = \frac{u_0 + v_0}{2}, \quad z = u_0 - v_0, \quad u_0 = w + \frac{z}{2}, \quad v_0 = w - \frac{z}{2}$$

$$\bar{T} = w - \frac{z}{2} + \eta z + P_2(\eta) \frac{Qs_2(\alpha_2 - \alpha_1)}{B+C}$$

On invariant 2D subspace: $z = z^* = \frac{Q(\alpha_2 - \alpha_1)}{B+C}$

$$\bar{T} = \bar{T}(w, \eta) = w + \left(\eta - \frac{1}{2}\right) \frac{Q(\alpha_2 - \alpha_1)}{B+C} + P_2(\eta) \frac{Qs_2(\alpha_2 - \alpha_1)}{B+C}$$

$$= w + \frac{Q(\alpha_2 - \alpha_1)}{B+C} \left(\eta - \frac{1}{2} + s_2 P_2(\eta)\right)$$


Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = \varepsilon \kappa (T_b(\eta, w) - T_c)$$


$$\dot{w} = \frac{\kappa}{R} \left(Q(1 - \alpha_0) - A + C\bar{T}(\eta, w) - (B+C)w - \varepsilon \Omega (T_b(\eta, w) - T_c) \right)$$

where

$$T_b(\eta, w) = w + \frac{Q(1 - \alpha_0)}{B+C} s_2 P_2(\eta)$$

$$\bar{T}(w, \eta) = w + \frac{Q(\alpha_2 - \alpha_1)}{B+C} \left(\eta - \frac{1}{2} + s_2 P_2(\eta) \right)$$

$$p_2(\eta) = \frac{1}{2}(3\eta^2 - 1)$$

$$P_2(\eta) = \frac{1}{2}(\eta^3 - \eta)$$


Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = \varepsilon \kappa (T_b(\eta, w) - T_c)$$

$$\dot{w} = \frac{\kappa}{R} \left(Q(1 - \alpha_0) - A + C\bar{T}(\eta, w) - (B+C)w - \varepsilon \Omega (T_b(\eta, w) - T_c) \right)$$

First, let $\varepsilon = 0$ and recall that


$$\bar{T}(w, \eta) = w + \frac{Q(\alpha_2 - \alpha_1)}{B+C} \left(\eta - \frac{1}{2} + s_2 P_2(\eta) \right)$$

$$\dot{w} = \frac{\kappa}{R} \left(Q(1 - \alpha_0) - A + C\bar{T}(\eta, w) - (B+C)w \right)$$

$$= \frac{\kappa}{R} \left(Q(1 - \alpha_0) - A - Bw + \frac{CQ(\alpha_2 - \alpha_1)}{B+C} \left(\eta - \frac{1}{2} + s_2 P_2(\eta) \right) \right)$$

$$= \frac{\kappa B}{R} (\Phi_0(\eta) - w)$$

where

$$\Phi_0(\eta) = \frac{1}{B} \left(Q(1 - \alpha_0) - A + \frac{CQ(\alpha_2 - \alpha_1)}{B+C} \left(\eta - \frac{1}{2} + s_2 P_2(\eta) \right) \right)$$


Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

For $\varepsilon = 0$, the system becomes


$$\dot{\eta} = 0$$

$$\dot{w} = \frac{\kappa B}{R} (\Phi_0(\eta) - w)$$

This system has a curve of fixed point given by

$$w = \Phi_0(\eta) = \frac{1}{B} \left(Q(1 - \alpha_0) - A + \frac{CQ(\alpha_2 - \alpha_1)}{B+C} \left(\eta - \frac{1}{2} + s_2 P_2(\eta) \right) \right)$$

This invariant curve is globally exponentially attracting with exponent

$$-\frac{\kappa B}{R}$$


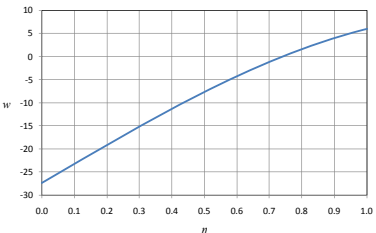

Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = 0$$

$$\dot{w} = \frac{\kappa B}{R} (\Phi_0(\eta) - w)$$

The curve of fixed points for $\varepsilon = 0$.

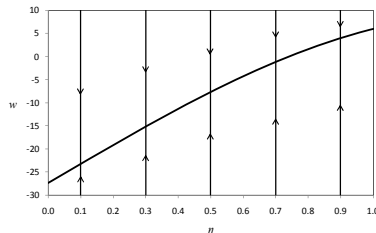
Ice-Albedo Feedback


Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = 0$$

$$\dot{w} = \frac{\kappa B}{R} (\Phi_0(\eta) - w)$$

Phase portrait for $\varepsilon = 0$.



 **Ice-Albedo Feedback**
Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = \varepsilon K \left(w + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c \right)$$

$$\dot{w} = \frac{K}{R} \left(B\Phi_0(\eta) - Bw - \varepsilon\Omega \left(w + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c \right) \right)$$


If we know η and w , what is the temperature profile $T(y)$?

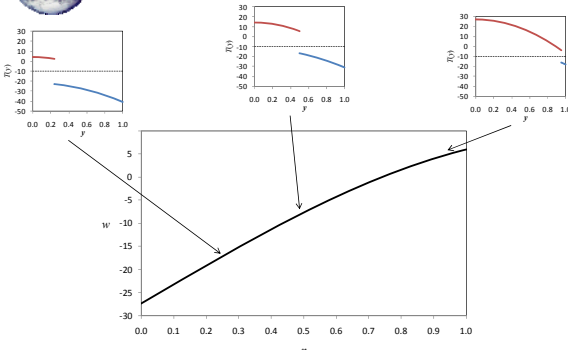
Recall that $T(y) = \begin{cases} U(y), & y < \eta, \\ V(y), & y > \eta, \end{cases}$ where $U(y) = u_0 + u_2 p_2(y)$, $V(y) = v_0 + v_2 p_2(y)$.


Recall also that $u_0 = w + \frac{\varepsilon}{2}$, $v_0 = w - \frac{\varepsilon}{2}$.

On the invariant 2-space $z = z^* = \frac{Q(\alpha_2 - \alpha_1)}{B+C}$, $u_2 = u_2^* = \frac{Qs_2(1-\alpha_1)}{(B+C)}$, $v_2 = v_2^* = \frac{Qs_2(1-\alpha_2)}{(B+C)}$.

Given η and w , we can compute the temperature profile $T(y)$.

 **Ice-Albedo Feedback**
Two Dimensional Budyko-Sellers-Widiasih Model



 **Ice-Albedo Feedback**
Two Dimensional Budyko-Sellers-Widiasih Model


$$\dot{\eta} = \varepsilon K \left(w + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c \right)$$

$$\dot{w} = \frac{K}{R} \left(B\Phi_0(\eta) - Bw - \varepsilon\Omega \left(w + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c \right) \right)$$

What about $\varepsilon > 0$?

Rest points: $w = -\frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) + T_c$, $w = \Phi_0(\eta)$

Solve for η : $h(\eta) = \Phi_0(\eta) + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c = 0$


 **Ice-Albedo Feedback**
Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = \varepsilon K \left(w + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c \right)$$

$$\dot{w} = \frac{K}{R} \left(B\Phi_0(\eta) - Bw - \varepsilon\Omega \left(w + \frac{Qs_2(1-\alpha_0)}{B+C} p_2(\eta) - T_c \right) \right)$$

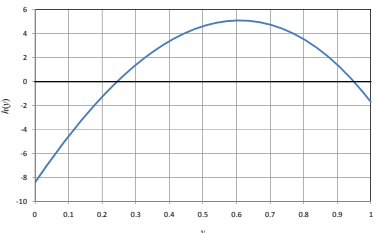
Note that rest points occur on the curve: $w = \Phi_0(\eta)$


and that near that curve, $\frac{d\eta}{dt} \approx \varepsilon K h(\eta)$

 **Ice-Albedo Feedback**
Two Dimensional Budyko-Sellers-Widiasih Model

$$\frac{d\eta}{dt} \approx \varepsilon K h(\eta)$$

The function h .

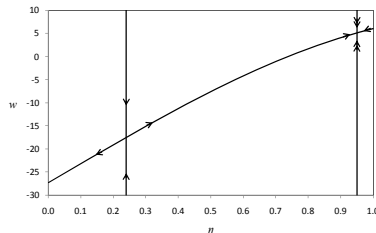


 **Ice-Albedo Feedback**
Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = \varepsilon K (T_s(\eta, w) - T_c)$$

$$\dot{w} = \frac{K}{R} (Q(1-\alpha_0) - A + C\bar{T}(\eta, w) - (B+C)w - \varepsilon\Omega(T_s(\eta, w) - T_c))$$

Phase portrait for tiny $\varepsilon > 0$.





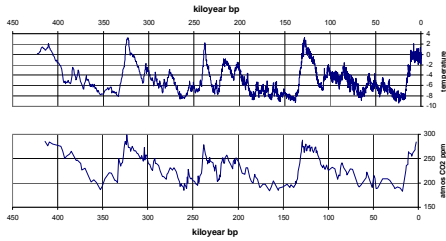
Ice-Albedo Feedback

Two Dimensional Budyko-Sellers-Widiasih Model

$$\dot{\eta} = \varepsilon \kappa (T_b(\eta, w) - T_c)$$

$$\dot{w} = \frac{\kappa}{R} (Q(1 - \alpha_0) - A + C\bar{T}(\eta, w) - (B + C)w - \varepsilon \Omega(T_b(\eta, w) - T_c))$$

Can the parameters be estimated from the ice core data?



Petit, et al, *Nature* 399 (June 3 1999), pp.429-436