


## Glacial Cycles: the 100,000-Year Problem


Richard McGehee



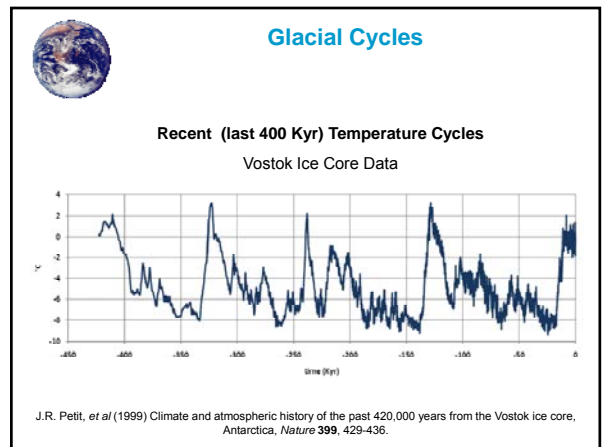
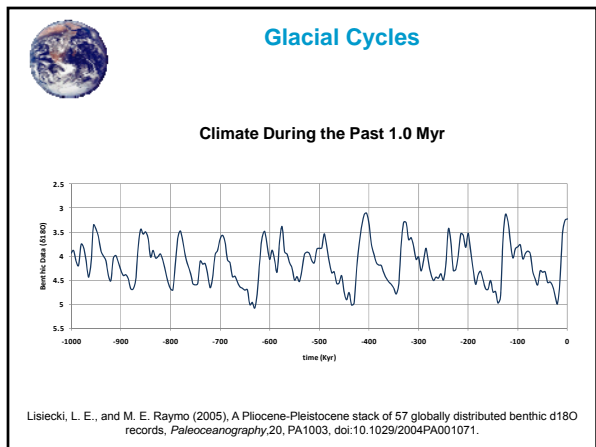
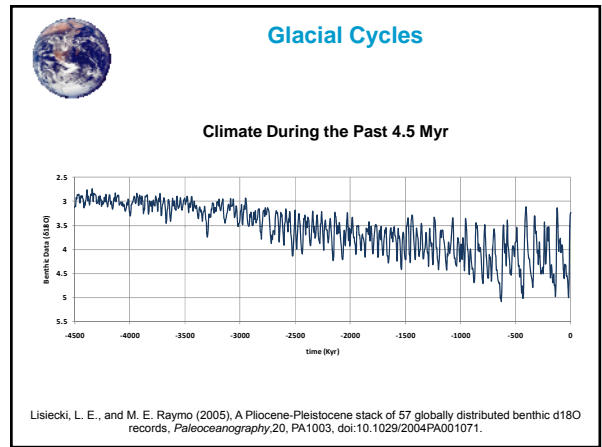
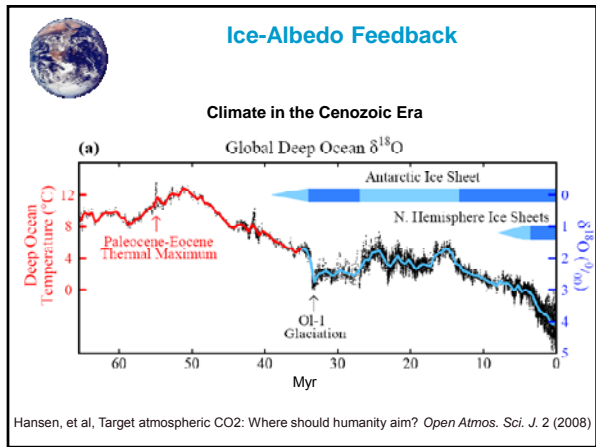
Seminar on the Mathematics of Climate Change  
School of Mathematics  
April 13, 2011

<http://www.tqmpc.org/NYC052141/beginningpage.html>

## Glacial Cycles: the 100,000-Year Problem



The 100,000-year problem is a discrepancy between past temperatures and the amount of incoming solar radiation, or insolation. The former rises and falls according to the strength of radiation from the sun, the distance from the earth to the sun, and the tilt of the Earth's poles. However, the ice-age cycle, which grows and shrinks periodically on a 100,000-year (100 ka) timescale, does not correlate well with any of these factors.





## Glacial Cycles

### What Causes Glacial Cycles?

#### Widely Accepted Hypothesis

The glacial cycles are driven by the variations in the Earth's orbit (Milankovitch Cycles), causing a variation in incoming solar radiation (insolation).

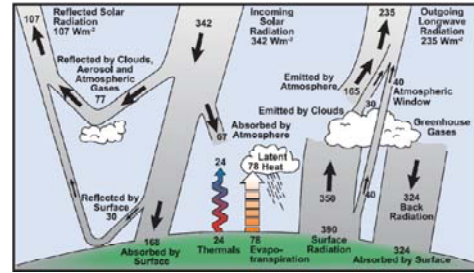
This hypothesis is widely accepted, but also widely regarded as insufficient to explain the observations.

The additional hypothesis is that there are feedback mechanisms that amplify the Milankovitch cycles. What these feedbacks are and how they work are not fully understood.



## Glacial Cycles

### Heat Balance

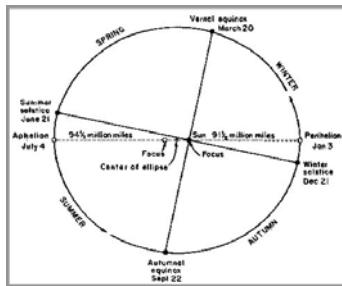


Historical Overview of Climate Change Science, IPCC AR4, p.96  
[http://ipcc-wg1.ucar.edu/wg1/Report/AR4WG1\\_Print\\_CH01.pdf](http://ipcc-wg1.ucar.edu/wg1/Report/AR4WG1_Print_CH01.pdf)



## Glacial Cycles

### Eccentricity

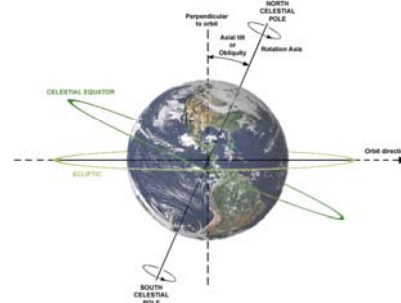


[http://www.crrel.usace.army.mil/permafrosttunnel/Ice\\_Age\\_Earth\\_Orbit.jpg](http://www.crrel.usace.army.mil/permafrosttunnel/Ice_Age_Earth_Orbit.jpg)



## Glacial Cycles

### Oblliquity

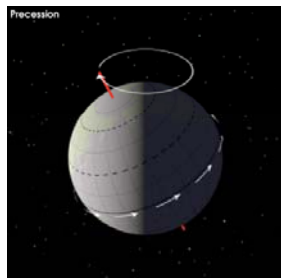


<http://upload.wikimedia.org/wikipedia/commons/6/61/AxialTiltobliquity.png>



## Glacial Cycles

### Precession

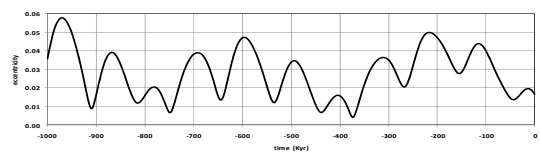


[http://earthobservatory.nasa.gov/Library/Giants/Milankovitch/milankovitch\\_2.html](http://earthobservatory.nasa.gov/Library/Giants/Milankovitch/milankovitch_2.html)



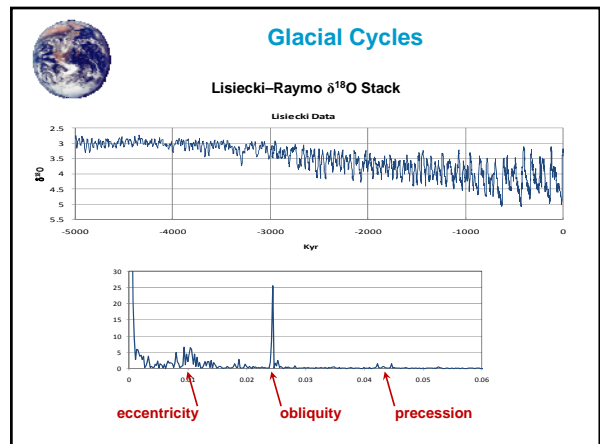
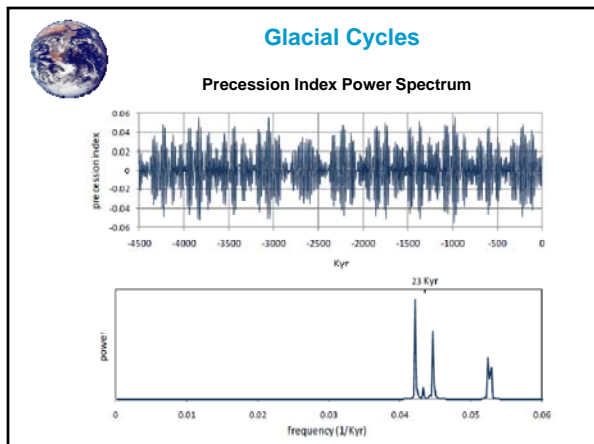
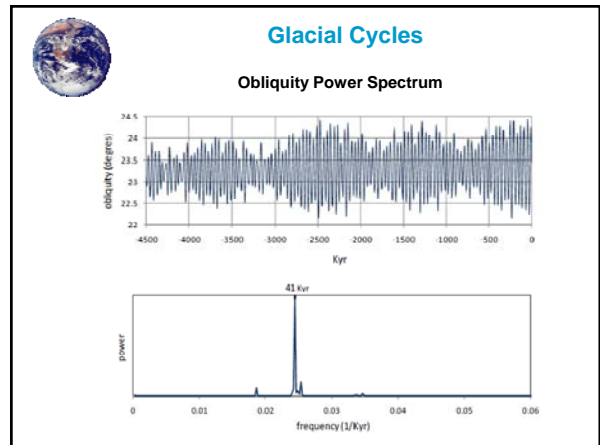
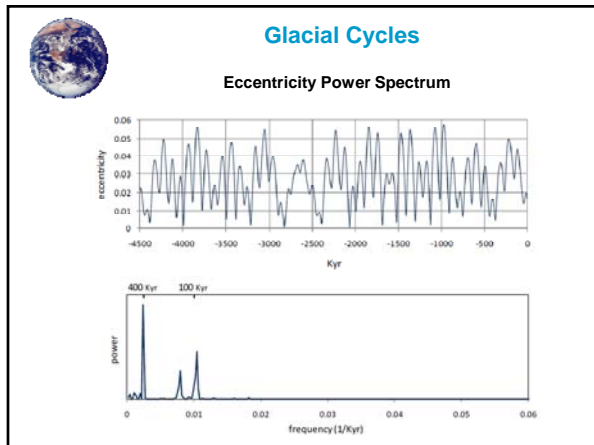
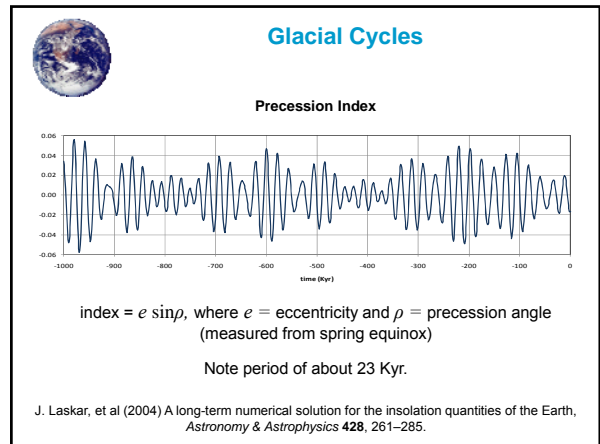
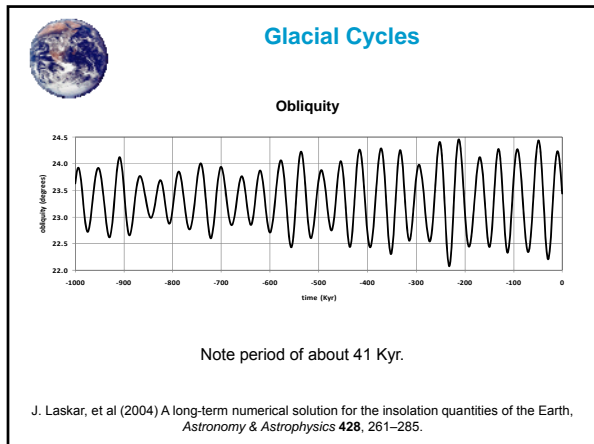
## Glacial Cycles

### Eccentricity



Note periods of about 100 Kyr and 400 Kyr.

J. Laskar, et al (2004) A long-term numerical solution for the insolation quantities of the Earth, *Astronomy & Astrophysics* 428, 261-285.



## Glacial Cycles

**Why obliquity and eccentricity?**

Incoming **Solar Radiation** (*Insolation*), averaged over the entire globe and over a full year, depends only on eccentricity  $e$ , not on either obliquity or precession.

$$Q(e) = \frac{Q_0}{\sqrt{1-e^2}}$$

Insolation as a function of latitude, averaged over a full year, depends on eccentricity  $e$  and obliquity  $\beta$ , but not precession.

$$I = Q(e)s(y, \beta)$$

where

$$s(y, \beta) = \frac{2}{\pi^2} \int_0^{2\pi} \sqrt{1 - (\sqrt{1-y^2} \sin \beta \cos \gamma - y \cos \beta)^2} d\gamma$$

$$y = \sin(\text{latitude})$$

## Glacial Cycles

**Why obliquity over eccentricity?**

Possible explanation: **Ice-albedo feedback**

Ice reflects more energy than land or water.  
 more ice → less energy → colder → more ice  
 less ice → more energy → warmer → less ice

## Glacial Cycles

**Heat Balance**

*Historical Overview of Climate Change Science, IPCC AR4, p.96  
[http://ipcc-wg1.ucar.edu/wg1/Report/AR4WG1\\_Print\\_CH01.pdf](http://ipcc-wg1.ucar.edu/wg1/Report/AR4WG1_Print_CH01.pdf)*

## Glacial Cycles

**Budyko-Sellers Model**

$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$

$\underbrace{Qs(y)}_{\text{insolation}} \underbrace{(1 - \alpha(y, \eta))}_{\text{albedo}} - \underbrace{(A + BT)}_{\text{re-radiation}} + \underbrace{C(\bar{T} - T)}_{\text{transport}}$

$T = T(y, t)$ : annual mean surface temperature  
 $y = \sin(\text{latitude}) \quad y \in [0, 1]$   
 $Q$ : global annual mean insolation  
 $s(t)$ : relative annual mean insolation  $\int_0^1 s(y) dy = 1$   
 $y = \eta$ : ice boundary  
 $\alpha(y, \eta) = \begin{cases} \alpha_1, & y < \eta \\ \alpha_2, & y > \eta \end{cases}$  albedo  
 $\bar{T}(t) = \int_0^1 T(y, t) dy$ : global annual mean temperature

## Glacial Cycles

Solve for equilibrium solution  $T^*(y)$ .  
 Set right hand side = 0.

equilibrium temperature profiles

Two equilibrium solutions:  
 small cap: stable  
 large cap: unstable

Not equilibria:  
 ice free  
 snowball

## Glacial Cycles

**Budyko-Sellers Model**

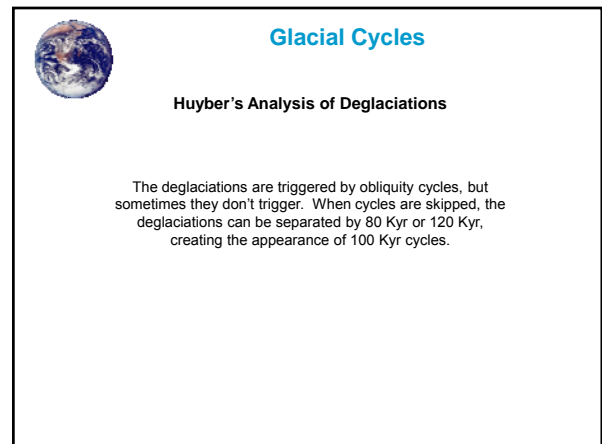
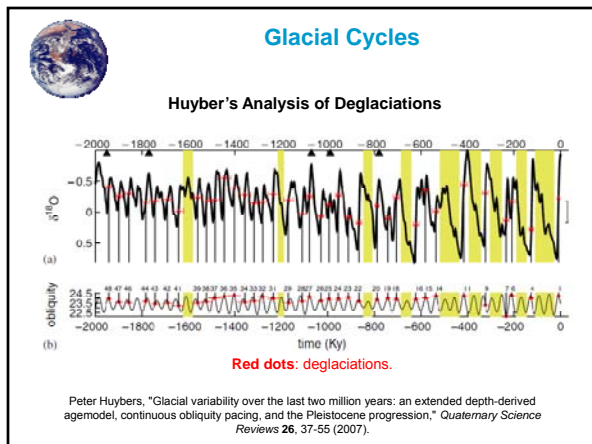
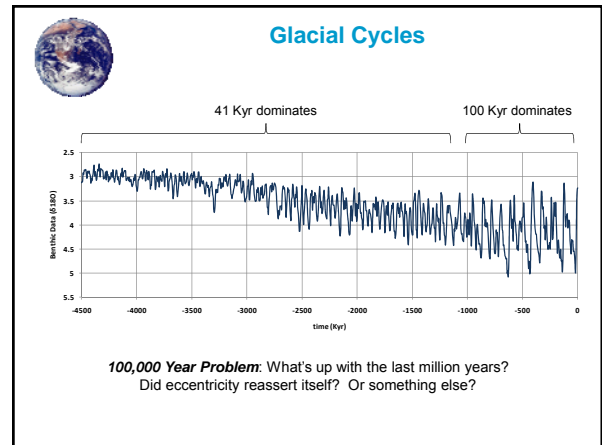
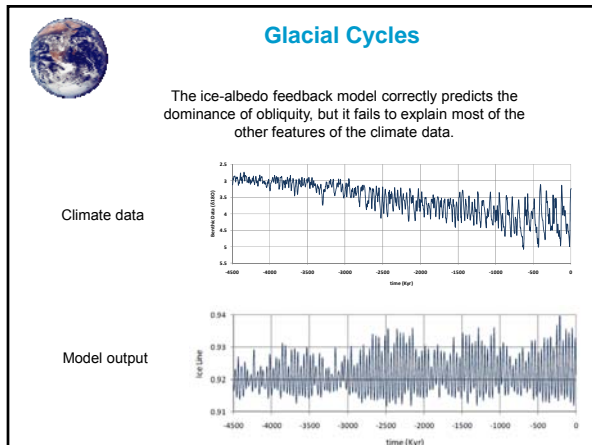
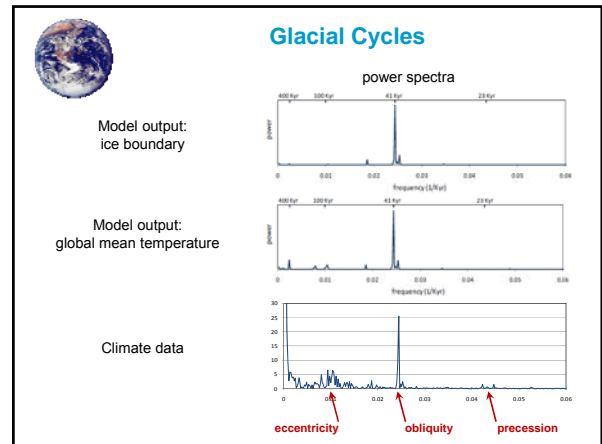
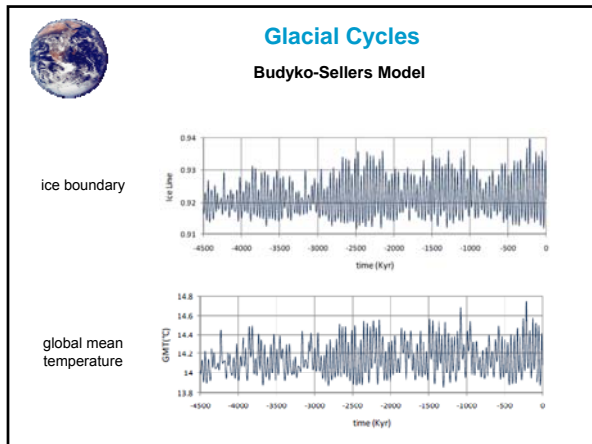
$$R \frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + C(\bar{T} - T)$$


Note that the equilibrium solution  $T^*(y)$  depends on  $Q$  and  $s(y)$ , which depend on the eccentricity  $e$  and the obliquity  $\beta$ . Therefore, the equilibrium location  $\eta$  of the ice boundary and the equilibrium global mean temperature (GMT) depend on the eccentricity and the obliquity.

We can use the computed values of eccentricity and obliquity to compute the ice boundary and GMT over the glacial cycles.

$$Q(e) = \frac{Q_0}{\sqrt{1-e^2}}$$

$$s(y, \beta) = \frac{2}{\pi^2} \int_0^{2\pi} \sqrt{1 - (\sqrt{1-y^2} \sin \beta \cos \gamma - y \cos \beta)^2} d\gamma$$





## Glacial Cycles

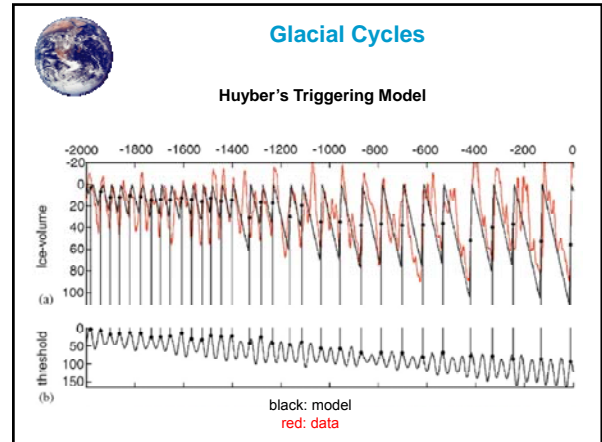

### Huyber's Triggering Model

$$V_t = \begin{cases} V_{t-1} + \eta & \text{if } V_t < T_t \\ 0 & \text{if } V_t \geq T_t \end{cases}$$

$V_t$ : ice volume at time  $t$   
 $T_t$ : threshold variable  
 $\eta$ : rate of increase of ice volume  
 $T_t = at + b - c\theta'_t$   
 $\theta'_t$ : normalized obliquity

*Units and constants*

$t$ : Kyr  
 $V$ : chosen so that  $\eta = 1$ .  
 $\theta'$ : mean zero and variance one  
 $a = 0.05$   
 $b = 126$   
 $c = 20$

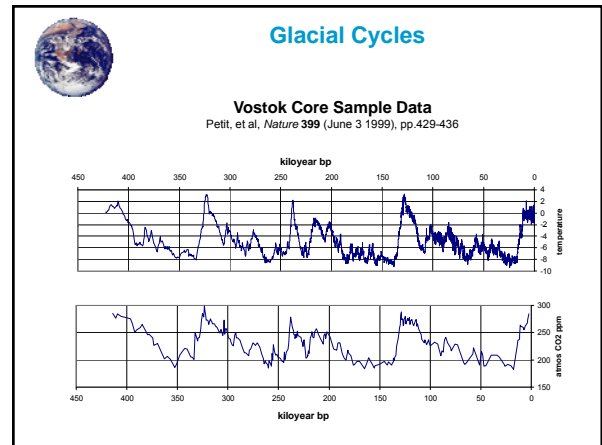

## Glacial Cycles

### Huyber's Triggering Model

Huyber's model produces the decline in temperature and the increase in period and amplitude of the glacial cycles, but it depends heavily on an unspecified decline in the sensitivity of the triggering mechanism over last two million years.

**What about greenhouse gases and the carbon cycle?**

Andrew Hogg suggested a model incorporating the carbon cycle.

## Glacial Cycles

### Hogg's Model

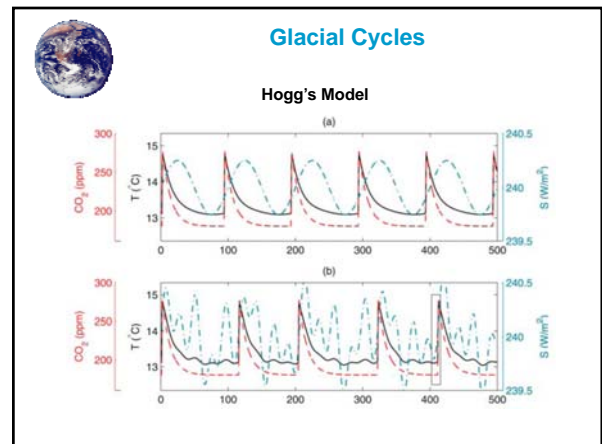
$$c \frac{dT}{dt} = S(t) + G(C) - \sigma T^4$$

$$\frac{dC}{dt} = V - \underbrace{(W_0 + W_1 C)}_{\text{weathering volcanos}} + \underbrace{\beta(C_{\text{max}} - C)}_{\text{CO2 outgassing}} \max\left(\frac{dT}{dt} - \epsilon, 0\right)$$

$$S(t) = \bar{S} + \sum_i S_i \sin\left(\frac{2\pi t}{\Gamma_i}\right) \quad \text{insolation}$$

$$G(C) = \bar{G} + A \ln\left(\frac{C}{C_0}\right) \quad \text{greenhouse forcing}$$

Andrew McC. Hogg, "Glacial cycles and carbon dioxide: A conceptual model," *Geophysical Research Letters* 35 (2008).





## Glacial Cycles

### Hogg's Model

Hogg's model shows how the carbon cycle can act as a feedback amplifying and modifying the insolation forcing, but the forcing is somewhat artificial, and the triggering mechanism is difficult to justify. Also, it does not solve the 100,000-year problem.

*What if the 100,000 year glacial cycle is not driven by eccentricity, but is a natural oscillation of the Earth's climate?*

Saltzman and Maasch suggested just such a model.



## Glacial Cycles

### Salzman-Maasch Model

$$\begin{aligned} \text{global ice mass} &\rightarrow \dot{X} = -X - Y - uM(t) \\ \text{atmospheric CO}_2 &\rightarrow \dot{Y} = -pZ + rY + sZ^2 - Z^2Y \\ \text{deep ocean temperature} &\rightarrow \dot{Z} = -q(X + Z) \end{aligned}$$

Milankovitch forcing

Barry Saltzman and Kirk A. Maasch, "A Low-Order Dynamical Model of Global Climatic Variability Over the Full Pleistocene," *Journal of Geophysical Research* 95 (D2), 1955-1963 (1990)



## Glacial Cycles

### Salzman-Maasch Model unforced

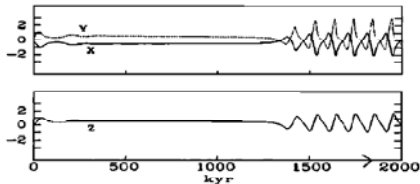


Fig. 7. Unforced (free) solution for (4)-(6) with  $q = 1.2$ ,  $s = 0.8$ , and  $p$  and  $r$  varying linearly between  $0.8 \rightarrow 1.0$  and  $0.7 \rightarrow 0.8$ , respectively. (Top) Model solutions for  $X$  (solid) and  $Y$  (dashed). (Bottom) Model solution for  $Z$ .



## Glacial Cycles

### Salzman-Maasch Model forced

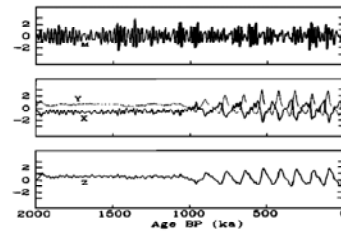


Fig. 8. Forced solution for (4)-(6) with  $q = 1.2$ ,  $s = 0.8$ ,  $u = 0.7$ , and  $p$  and  $r$  varying linearly between  $0.8 \rightarrow 1.0$  and  $0.7 \rightarrow 0.8$ , respectively. (Top) Normalized  $65^\circ\text{N}$  July insolation curve ( $M$ ) used as forcing. (Middle) Model solutions for  $X$  (solid) and  $Y$  (dashed). (Bottom) Model solution for  $Z$ .



## Glacial Cycles

### Salzman-Maasch Model

The Salzman-Maasch model shows how the carbon cycle and the ocean currents can interact to produce unforced oscillations with periods of about 100,000 years. The same model with slightly different parameters can exhibit stationary behavior. By forcing the model with Milankovitch cycles and by slowly varying the parameters over the last two million years, they can produce a bifurcation from small oscillations tracking the Milankovitch cycles to large oscillations with a dominant 100,000 year period.

*Seems like a nice idea, but it is not widely accepted as the explanation.*



## Glacial Cycles

### Current Project

The Mathematics and Climate Research Network (MCRN) has a Webinar working group developing a model incorporating ice-albedo feedback with the carbon cycle.

*Local expert: Samantha Oestriecher*



## Glacial Cycles

### The 100,000-Year Problem

#### Summary

100 Kyr cycles during the last million years, but 41 Kyr cycles before that.

#### **Why?**

**Huybers:** Obliquity rules, but glaciers started skipping beats.  
Alternating 80 Kyr and 120 Kyr looks like 100 Kyr

**Saltzman & Maasch:** Under some conditions, the climate naturally oscillates at 100 Kyr. Those conditions arose 1 Myr ago. Before that, the climate tracked Milankovitch.

#### **Other ... ?**