

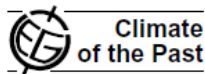
# Ecosystem effects of CO<sub>2</sub> concentration: a paper review on how C<sub>3</sub> & C<sub>4</sub> plants interact with CO<sub>2</sub> concentration in paleodata

Samantha Oestreicher  
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## Ecosystem effects of CO<sub>2</sub> concentration: evidence from past climates

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# Three types of Photosynthesis: $C_3$ , $C_4$ , and CAM

## 1 $C_3$

- ▶ Most plants are  $C_3$ . Photosynthesis takes place throughout the leaf.
- ▶ More efficient than  $C_4$  or CAM in moist, cool conditions because photosynthesis requires less machinery.
- ▶ Atmospheric  $CO_2$  is the limiting factor for photosynthesis even for today's elevated levels.
- ▶ **Examples:** wheat, barley, potatoes and sugar beet.



# Three types of Photosynthesis: C<sub>3</sub>, C<sub>4</sub>, and CAM

## 2 C<sub>4</sub>

- ▶ Photosynthesis takes place in inner cells.
- ▶ Photosynthesises faster than C<sub>3</sub>
- ▶ Adapted for low atmospheric CO<sub>2</sub>.
- ▶ **Examples:** fourwing saltbush, corn, many of summer annual plants.



# Three types of Photosynthesis: $C_3$ , $C_4$ , and CAM

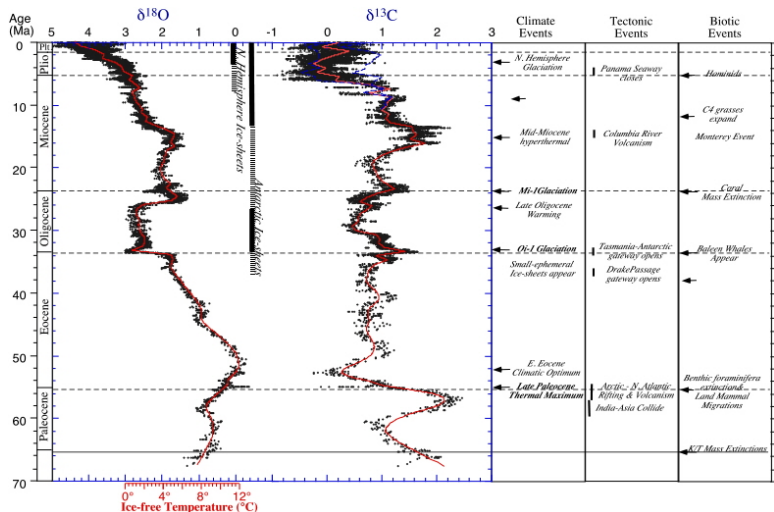
## 3 CAM

- ▶ Photosynthesis has diurnal cycle.
- ▶ Adapted for very arid environments.
- ▶ **Examples:** cactuses, some orchids and bromeliads



<http://legacy.earlham.edu/vandeel/notes.htm>

# Affecting Paleoclimate Data



## Affecting Paleoclimate Data

Atmospheric CO<sub>2</sub> concentration,  $c_a$ , affects C<sub>3</sub> photosynthesis and therefore will effect competition between C<sub>3</sub> and C<sub>4</sub> plants.

One example of this could be between C<sub>3</sub> trees and C<sub>4</sub> grasses in a tropical savannah. This could affect pollen assemblages!



## Affecting Paleoclimate Data

“If these variations in  $c_a$  have caused changes that are detectable in compositional data, such as pollen assemblages, then conventional approaches to reconstructing past climate using statistical or analogue methods - if applied to period with  $c_a$  different from that of the late Holocene- are certain to yield incorrect results.”



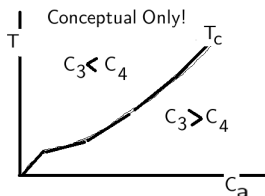
## Affecting Paleoclimate Data

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I think this means that if we assume the same ratio of  $C_3$  to  $C_4$  plants as we have today when we do pollen reconstruction, then we'll get the wrong answer.

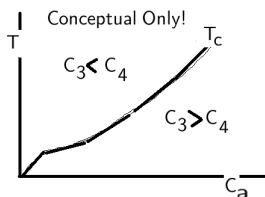
## C<sub>3</sub> vs C<sub>4</sub> Carbon fixing

- There exists a crossover temperature  $T_c$  above which C<sub>4</sub> plants can fix CO<sub>2</sub> much faster than C<sub>3</sub> plants.
- $T_c$  increases as  $c_a$  increases.
- $T < T_c \Rightarrow$  that C<sub>4</sub> is fixing carbon much slower. Thus C<sub>4</sub> is more competitive at low  $c_a$ .
- $\therefore$  today we should expect C<sub>3</sub> plants to be dominating.



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*A widespread trend towards an increase of (C<sub>3</sub>) tree cover at the expense of (C<sub>4</sub>) grasses has indeed been observed in tropical savannahs.*

# Understanding Glacial Terrestrial Carbon Storage Decrease

- In 1977, Shackleton notes that subfossil benthic foraminifera from LGM were light in  $^{13}\text{C}$ .
- This suggests a change in the stable isotope composition in the ocean.
- Relative to the Holocene,  $\delta^{13}\text{C}$  was offset by  $-0.7$  per mille at LGM
- Shackleton's hypothesis is that this change is caused by a balance between terrestrial and ocean carbon reservoirs because it is only terrestrial organic carbon which is observed to be depleted in  $^{13}\text{C}$ .
- A simple mass balance says organic carbon storage was greatly reduced during glacial times.

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Can we put a number on “greatly”?

He wasn't quite correct. The canonical estimate is  $-0.32$  per mille offset in  $\delta^{13}\text{C}$  from benthic data.

# Estimates of Terrestrial Carbon Storage Reduction for LGM

310 – 550 <i>PgC</i>	Bird et al(1994, 1996)
300 – 700 <i>PgC</i>	Bird et al(1996)
430 – 665 <i>PgC</i>	Strecet – Perrott et al(1998)
550 – 680 <i>PgC</i>	Beerling(1999)
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Two interesting comments Prentice et al makes:

(1) Reduced carbon storage is **not** accounted for by the presense of continental ice sheets becuase the area of exposed continental shelf (in the tropics) roughly balances that occupied by ice!

(2) Peatland **cannot** account for “net postglacial carbon accumulation” becuase there is evidence of tropic peatlands which sequester on the order of 200 PgC.

## Physiological Effects

Physiological Effects (physical plant mechanism differences) can reduce carbon storage on land during a glacial maximum.

$C_4$  plants fix carbon much slower than  $C_3$ .  $C_4$  plants are more dominant in a glacial maximum. Therefore less carbon is being sequestered per unit of time.



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However, there is a opposing mechanism. Carbon stays in the soil much longer at low temperatures because decomposition happens much slower. (hence all the really good soil is in Boreal regions).

*One model suggests that global cooling to LGM levels would have a minor impact on total terrestrial carbon storage, wherea the  $CO_2$  changes would reduce terrestrial carbon storage substantially.*

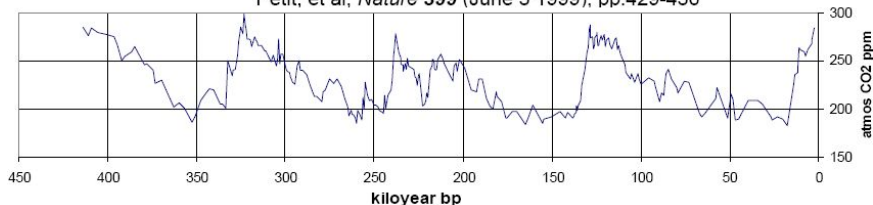
# The Claim

Paleodata combines with process based modelling that includes physiological effects can correctly reproduce

- 1 The 100ppm increase in  $c_a$  from LGM to Holocene.
- 2 The approx. magnitude of carbon storage change.

## Vostok Core Sample Data

Petit, et al, *Nature* 399 (June 3 1999), pp.429-436



# Determining Magnitude of Carbon Storage

- View pollen and plant macro-fossil records from terrestrial sediments with an eye for change in  $c_a$ .
- Many papers have shown that  $\text{CO}_2$  changes with change plant types, biomes etc but haven't added it to the model before.
- There exists evidence that tropical lowland regions shift towards  $\text{C}_4$  plants during glacial times.

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**Question:** How do we determine if Physiological Effects are important?

**Answer:** Run some really big models with and without the code that determines physiological effects.

A typical result is that the terrestrial carbon storage reduction is:

160 PgC without physiological effects.

610 PgC with physiological effects.

## Conclusion

Physiological effects are essential to correctly modeling terrestrial carbon storage changes.

Additionally, if one wants to model previous climates, then one must consider the competition between  $C_3$  and  $C_4$  plant to determine the  $CO_2$  fixing values and other plant type parameters.

