


**Math 5490**  
**Topics in Applied Mathematics**  
**Introduction to the Mathematics of Climate**


Fall 2023  
**1:25 - 3:20 Tuesdays and Thursdays**  
**Amundson Hall 162**

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course website  
[www-users.cse.umn.edu/~mcgehee/teaching/Math5490/](http://www-users.cse.umn.edu/~mcgehee/teaching/Math5490/)



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**Ocean Circulation**


**Welander's Model**

**Reference**

Pierre Welander, A simple heat-salt oscillator,  
*Dynamics of Atmospheres and Oceans* 6 (1982)  
 233-242.

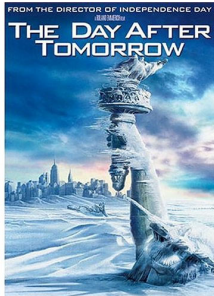


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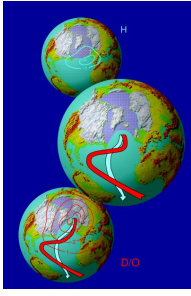
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**Heinrich and Dansgaard-Oeschger events**




FROM THE DIRECTOR OF INDEPENDENCE DAY  
**THE DAY AFTER TOMORROW**


What did the film get right scientifically?




<http://www.pik-potsdam.de/~stefan/sampleimages.html>



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 **INDEPENDENT**  
 April 12, 2018


**Gulf Stream current at 'record low' with potentially devastating consequences for weather, warn scientists**

The Atlantic meridional overturning circulation (AMOC), the system of currents that transports warm water from the tropics via the Gulf Stream to the North Atlantic, plays a major role in regulating the world's climate.

A fictional depiction of AMOC's collapse was portrayed in *The Day After Tomorrow*, and while the film's events were exaggerated, scientists say severe weather events are likely to result from the ongoing changes.



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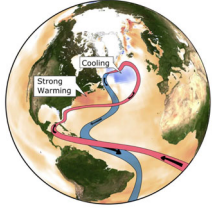
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**The Washington Post**  
 April 11, 2018

**The oceans' circulation hasn't been this sluggish in 1,000 years. That's bad news.**


The Atlantic Ocean circulation that carries warmth into the Northern Hemisphere's high latitudes is slowing down because of climate change, a team of scientists asserted Wednesday, suggesting one of the most feared consequences is already coming to pass.

[Nature](#) volume 556, pages191–196 (2018)





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**Current Atlantic Meridional Overturning Circulation weakest in last millennium**

The Atlantic Meridional Overturning Circulation (AMOC)—one of Earth's major ocean circulation systems—redistributes heat on our planet and has a major impact on climate. Here, we compare a variety of published proxy records to reconstruct the evolution of the AMOC since about AD 400. A fairly consistent picture of the AMOC emerges: **after a long and relatively stable period, there was an initial weakening starting in the nineteenth century, followed by a second, more rapid, decline in the mid-twentieth century, leading to the weakest state of the AMOC occurring in recent decades.**

NATURE GEOSCIENCE | VOL 14 | MARCH 2021 | 118–120 | [www.nature.com/naturegeoscience118](http://www.nature.com/naturegeoscience118)

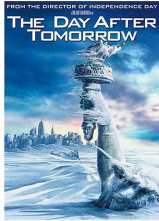


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**The Day After Tomorrow/Film synopsis**

After climatologist Jack Hall (Dennis Quaid) is largely ignored by U.N. officials when presenting his environmental concerns, his research proves true when an enormous "superstorm" develops, setting off catastrophic natural disasters throughout the world. Trying to get to his son, Sam (Jake Gyllenhaal), who is trapped in New York with his friend Laura (Emmy Rossum) and others, Jack and his crew must travel by foot from Philadelphia, braving the elements, to get to Sam before it's too late.



Google Search

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**Dansgaard-Oeschger Events**

"Global warming" can cause the Northern Hemisphere to cool.

Melting ice can lower the salinity of the North Atlantic, causing a decrease in the flow of the Atlantic Meridional Overturning Circulation (AMOC), slowing the heat transfer to the Northern Hemisphere.

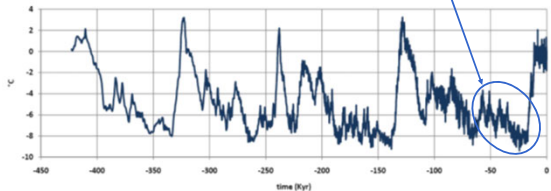
This phenomenon is believed to have caused the Younger Dryas.



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**Recent (last 400 Kyr) Temperature Cycles**  
**Vostok Ice Core Data**



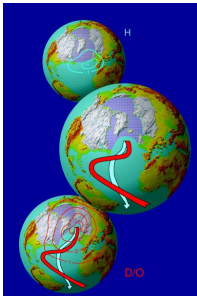
*What's with these oscillations?*

J.R. Petit, et al (1999) Climate and atmospheric history of the past 420,000 years from the Vostok ice core, Antarctica, *Nature* 399, 429-436.

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**Heinrich and Dansgaard-Oeschger events**

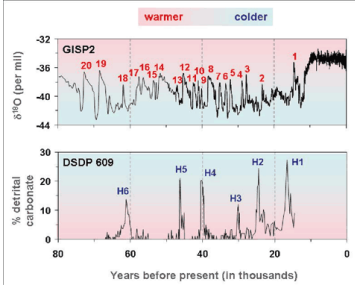


<http://www.pik-potsdam.de/~stefan/sampleimages.html>

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**Heinrich and Dansgaard-Oeschger events**

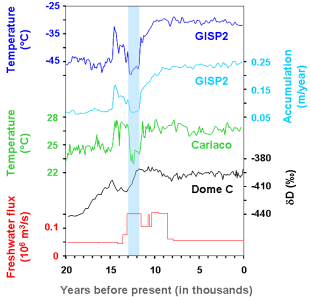


<http://www.ncdc.noaa.gov/paleo/abrupt/data3.html>


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**The Younger Dryas**

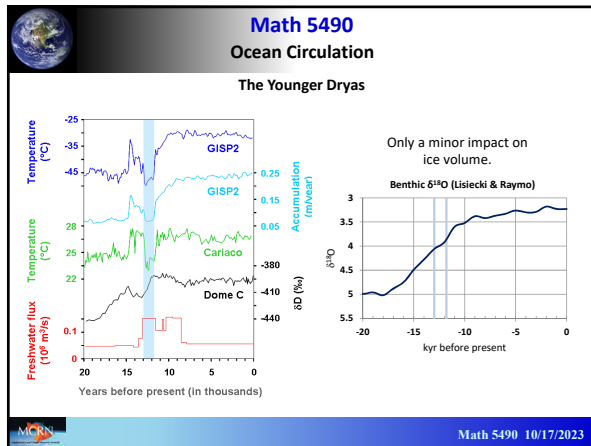


**Mountain Avens (*Dryas octopetala*)**



<https://www.ncdc.noaa.gov/abrupt-climate-change/the120Younger120dryas>

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**What caused the Dansgaard-Oeschger Oscillations?**

For glacial cycles, there is overwhelming evidence that they are "paced" by the Milankovitch cycles.

None of the Milankovitch cycles has periods short enough to trigger the Dansgaard-Oeschger events, and we can't think of any other external triggers.

Could they occur spontaneously?

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**What caused the Dansgaard-Oeschger oscillations?**

They could be self-oscillations in the natural dynamics of ocean circulation.

Pierre Welander, A simple heat-salt oscillator, *Dynamics of Atmospheres and Oceans* 6 (1982) 233-242.

R/V Weelander is a 23-foot-long Beach Master work boat, informally named in honor of Professor Pierre Welander (1925–1996).

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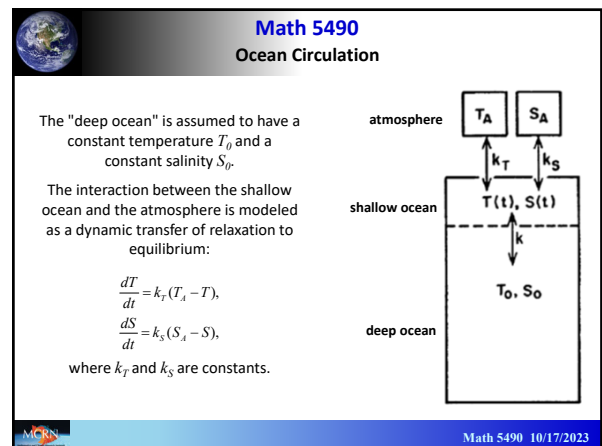
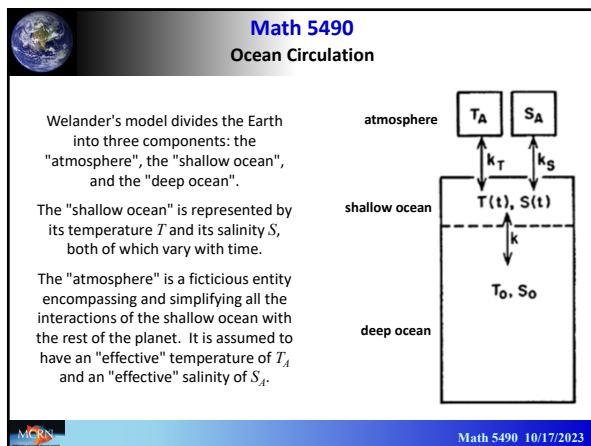
**What caused the Dansgaard-Oeschger oscillations?**

They could be self-oscillations in the natural dynamics of ocean circulation.

Welander constructed a simple (*conceptual*) box model of ocean circulation and showed that the interactions of temperature and salinity with the atmosphere, the surface ocean, and the deep ocean could create self-oscillations.

R/V Weelander is a 23-foot-long Beach Master work boat, informally named in honor of Professor Pierre Welander (1925–1996).

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**Math 5490**  
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**Discussion**  
How do we solve this differential equation?

$$\frac{dT}{dt} = k_r(T_a - T)$$

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**Math 5490**  
Ocean Circulation

**Discussion**  
How do we solve this differential equation?

"separate variables"  $\frac{dT}{dt} = k_r(T_a - T)$

$$\frac{dT}{T_a - T} = k_r dt \quad \frac{dT}{T - T_a} = -k_r dt \quad \int \frac{dT}{T - T_a} = \int -k_r dt$$

$$\ln(T - T_a) = -k_r t + c \quad \leftarrow \text{arbitrary constants}$$

$$T - T_a = e^{-k_r t + c} = e^c e^{-k_r t} = C e^{-k_r t}$$

$$T(t) = T_a + C e^{-k_r t}$$

**general solution**

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Ocean Circulation

**Discussion**  
How do we solve this differential equation?

$$\frac{dT}{dt} = k_r(T_a - T)$$

$$T(t) = T_a + C e^{-k_r t}$$

**general solution**

**check**

$$\frac{dT}{dt} = \frac{d}{dt}(T_a + C e^{-k_r t}) = 0 + C e^{-k_r t} (-k_r) = -k_r C e^{-k_r t}$$

$$k_r(T_a - T) = k_r(T_a - (T_a + C e^{-k_r t})) = -k_r C e^{-k_r t} \quad \leftarrow \text{checks!}$$

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**Discussion**

$$\frac{dT}{dt} = k_r(T_a - T)$$

**Find an equilibrium solution.**

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**Discussion**

$$\frac{dT}{dt} = k_r(T_a - T)$$

**Find an equilibrium solution.**

**Answer**

$$0 = \frac{dT}{dt} = k_r(T_a - T)$$

$$T(t) = T_a$$

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**Discussion**

$$\frac{dT}{dt} = k_r(T_a - T)$$

**equilibrium solution:**  $T(t) = T_a$

**Is it stable?**

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**Discussion**

$$\frac{dT}{dt} = k_T(T_A - T)$$

equilibrium solution:  $T(t) = T_A$

Is it stable?

Yes!

general solution:  $T(t) = T_A + Ce^{-k_T t}$

$$\lim_{t \rightarrow \infty} T(t) = T_A$$

Note that we are assuming that  $k_T$  is positive.

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**Discussion**

How do we solve this "initial value problem"?

$$\frac{dT}{dt} = k_T(T_A - T), \quad \text{equation}$$

$$T = T_0 \text{ when } t = 0. \quad \text{initial value}$$

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**Discussion**

How do we solve this "initial value problem"?

$$\frac{dT}{dt} = k_T(T_A - T),$$

$$T = T_0 \text{ when } t = 0.$$

general solution:  $T(t) = T_A + Ce^{-k_T t}$

$$T(0) = T_A + Ce^{-k_T \cdot 0} = T_A + C = T_0$$

$$C = T_0 - T_A$$

$$T(t) = T_A + Ce^{-k_T t}$$

$$T(t) = T_A + (T_0 - T_A)e^{-k_T t}$$

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**Discussion**

How do we solve this "initial value problem"?

$$\frac{dT}{dt} = k_T(T_A - T),$$

$$T = T_0 \text{ when } t = 0.$$

solution:  $T(t) = T_A + (T_0 - T_A)e^{-k_T t}$

check:

$$T'(t) = -k_T(T_0 - T_A)e^{-k_T t}$$

$$k_T(T_A - T) = k_T(-(T_0 - T_A)e^{-k_T t})$$

equal

$$T(0) = T_A + (T_0 - T_A)e^{-k_T \cdot 0} = T_A + (T_0 - T_A) = T_0$$

checks!

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**Discussion**

How do we solve this "initial value problem"?

$$\frac{dT}{dt} = k_T(T_A - T), \quad \text{equation}$$

$$T = T_0 \text{ when } t = 0. \quad \text{initial value}$$

alternative approach:

Introduce the departure from equilibrium

$$y = T - T_A, \quad T = T_A + y$$

$$\frac{dy}{dt} = \frac{dT}{dt} = k_T(T_A - T) = k_T(-y) = -k_T y$$

new initial value problem

$$\frac{dy}{dt} = -k_T y,$$

$$y = T_0 - T_A \text{ when } t = 0.$$

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**Discussion**

How do we solve this "initial value problem"?

$$\frac{dy}{dt} = -k_T y,$$

$$y = T_0 - T_A \text{ when } t = 0.$$

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**Discussion**  
How do we solve this "initial value problem"?

"separate variables"  
 $\frac{dy}{dt} = -k_T y,$   
 $y = T_0 - T_A$  when  $t = 0.$

$\frac{dy}{y} = -k_T dt$      $\frac{dy}{y} = -k_T dt$      $\int_{T_0 - T_A}^y \frac{dy}{y} = \int_0^t -k_T dt$   
 $\ln y \Big|_{T_0 - T_A}^y = -k_T t \Big|_0^t$      $\ln y - \ln(T_0 - T_A) = -k_T t$      $\frac{y}{T_0 - T_A} = e^{-k_T t}$

$y(t) = (T_0 - T_A) e^{-k_T t}$   
**stable**  
 $y \rightarrow 0$  as  $t \rightarrow 0$

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**Summary**

$\frac{dT}{dt} = k_T(T_A - T)$

This differential equation has a stable equilibrium solution  $T = T_A$ .

**Remember Welander?**

$T$  is the Welander temperature.

Salinity:  
 $\frac{dS}{dt} = k_S(S_A - S)$

This differential equation also has a stable equilibrium solution  $S = S_A$ .

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**Welander's Model**

The interaction between the shallow ocean and the atmosphere is modeled as a dynamic transfer of relaxation to equilibrium:

$\frac{dT}{dt} = k_T(T_A - T),$   
 $\frac{dS}{dt} = k_S(S_A - S),$

where  $k_T$  and  $k_S$  are positive constants. This system has a stable equilibrium point at  $(T, S) = (T_A, S_A)$ .

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**Welander's Model**

**Atmosphere - Ocean Surface Interaction**

system of differential equations

$\left\{ \begin{aligned} \frac{dT}{dt} &= k_T(T_A - T) = -k_T(T - T_A), & T(0) &= T_0, \\ \frac{dS}{dt} &= k_S(S_A - S) = -k_S(S - S_A), & S(0) &= S_0. \end{aligned} \right.$  initial condition

**Matrix Notation**

differential equation

$\frac{d}{dt} \begin{bmatrix} T \\ S \end{bmatrix} = \begin{bmatrix} -k_T & 0 \\ 0 & -k_S \end{bmatrix} \begin{bmatrix} T - T_A \\ S - S_A \end{bmatrix}, \quad \begin{bmatrix} T \\ S \end{bmatrix}(0) = \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}$  initial condition

**Solution**

singular! ONE vector equation

$\begin{bmatrix} T \\ S \end{bmatrix}(t) = \begin{bmatrix} T(t) \\ S(t) \end{bmatrix} = \begin{bmatrix} T_A + (T_0 - T_A)e^{-k_T t} \\ S_A + (S_0 - S_A)e^{-k_S t} \end{bmatrix}$

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**Dynamical Systems**

AMERICAN MATHEMATICAL SOCIETY  
COLLEQUIUM PUBLICATIONS, VOLUME 18  
DYNAMICAL SYSTEMS  
BY  
GEORGE D. BIRKHOFF, PH.D., M.S.  
PROFESSOR OF MATHEMATICS  
HARVARD UNIVERSITY  
PROVIDENCE  
PUBLISHED BY THE  
AMERICAN MATHEMATICAL SOCIETY  
1927

[https://en.wikipedia.org/wiki/Henri\\_Poincar%C3%A9](https://en.wikipedia.org/wiki/Henri_Poincar%C3%A9)

[https://en.wikipedia.org/wiki/George\\_David\\_Birkhoff](https://en.wikipedia.org/wiki/George_David_Birkhoff)

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**Dynamical Systems**

**Basic idea:**  
The dependence of the solution on **initial conditions** is just as important as its dependence on **time**.

**Initial value problem**

differential equation

$\frac{d}{dt} \begin{bmatrix} T \\ S \end{bmatrix} = \begin{bmatrix} -k_T & 0 \\ 0 & -k_S \end{bmatrix} \begin{bmatrix} T - T_A \\ S - S_A \end{bmatrix}, \quad \begin{bmatrix} T \\ S \end{bmatrix}(0) = \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}$

**solution**

$\begin{bmatrix} T \\ S \end{bmatrix}(t) = \begin{bmatrix} T(t) \\ S(t) \end{bmatrix} = \begin{bmatrix} T_A + (T_0 - T_A)e^{-k_T t} \\ S_A + (S_0 - S_A)e^{-k_S t} \end{bmatrix}$

**flow**

$\phi \left( \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}, t \right) = \begin{bmatrix} T_A + (T_0 - T_A)e^{-k_T t} \\ S_A + (S_0 - S_A)e^{-k_S t} \end{bmatrix}$

initial condition    time

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**Ocean Circulation**

**Dynamical Systems**

**Basic idea:**  
The dependence of the solution on *initial conditions* is just as important as its dependence on *time*.

**General notation:**

state variable  $x = \begin{bmatrix} T \\ S \end{bmatrix}$     vector field  $f(x) = \begin{bmatrix} -k_T & 0 \\ 0 & -k_S \end{bmatrix} \begin{bmatrix} T - T_A \\ S - S_A \end{bmatrix}$     initial condition  $\xi = \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}$     dot notation  $\dot{x} = \frac{dx}{dt}$     differential equation  $\dot{x} = f(x)$

**flow**  
 $\varphi(\xi, t) = \varphi \left( \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}, t \right) = \begin{bmatrix} T_A + (T_0 - T_A)e^{-k_T t} \\ S_A + (S_0 - S_A)e^{-k_S t} \end{bmatrix}$

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**Dynamical Systems**

The dependence of the solution on *initial conditions* is just as important as its dependence on *time*.

$x \in \mathbb{R}^n, \xi \in \mathbb{R}^n, f: \mathbb{R}^n \rightarrow \mathbb{R}^n$

$\dot{x} = f(x)$   
 $x(0) = \xi$

**initial value problem**

The initial value problem generates a flow  $\varphi: \mathbb{R}^n \times \mathbb{R} \rightarrow \mathbb{R}^n$  with properties

**initial condition**  $\varphi(\xi, 0) = \xi$

**"group property"**  $\varphi(\varphi(\xi, t), s) = \varphi(\xi, t+s)$

$\varphi(\eta, s) = \varphi(\xi, t+s)$   
 $\eta = \varphi(\xi, t)$

If we start the system at state  $\xi$  and follow the solution for time  $t$ , then restart the system at the new state and follow the solution for time  $s$ , we end up at the same state as starting at  $\xi$  and following for time  $t+s$ .

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**Dynamical Systems**

**Example**

**initial value problem**  
 $\frac{d}{dt} \begin{bmatrix} T \\ S \end{bmatrix} = \begin{bmatrix} -k_T & 0 \\ 0 & -k_S \end{bmatrix} \begin{bmatrix} T - T_A \\ S - S_A \end{bmatrix}$   
 $\begin{bmatrix} T \\ S \end{bmatrix}(0) = \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}$

**flow**  
 $\varphi \left( \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}, t \right) = \begin{bmatrix} T_A + (T_0 - T_A)e^{-k_T t} \\ S_A + (S_0 - S_A)e^{-k_S t} \end{bmatrix}$

**group property**  
 $\varphi \left( \varphi \left( \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}, t \right), s \right) = \begin{bmatrix} T_A + (T_0 - T_A)e^{-k_T(t+s)} \\ S_A + (S_0 - S_A)e^{-k_S(t+s)} \end{bmatrix} = \varphi \left( \begin{bmatrix} T_0 \\ S_0 \end{bmatrix}, t+s \right)$

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**Dynamical Systems**

**Example**  
 $\frac{dx}{dt} = \alpha x$

$\alpha < 0$      $\alpha = 0$      $\alpha > 0$

"asymptotically stable"    "Lyapunov stable"    "unstable"

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**Dynamical Systems**

**"Phase Plane"**

**Example**

$\frac{d}{dt} \begin{bmatrix} x \\ y \end{bmatrix} = \begin{bmatrix} -1 & 0 \\ 0 & -1 \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}$      $\dot{x} = -x$      $\dot{y} = -y$

$\begin{bmatrix} x \\ y \end{bmatrix}(0) = \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}$      $x(0) = x_0$      $y(0) = y_0$

$\varphi \left( \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}, t \right) = \begin{bmatrix} x_0 e^{-t} \\ y_0 e^{-t} \end{bmatrix}$

Phase plane plot showing trajectories converging to the origin from all directions.

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**Dynamical Systems**

**"Phase Plane"**

**Example**

$\frac{d}{dt} \begin{bmatrix} x \\ y \end{bmatrix} = \begin{bmatrix} -1 & 0 \\ 0 & -2 \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}$      $\dot{x} = -x$      $\dot{y} = -2y$

$\begin{bmatrix} x \\ y \end{bmatrix}(0) = \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}$      $x(0) = x_0$      $y(0) = y_0$

$\varphi \left( \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}, t \right) = \begin{bmatrix} x_0 e^{-t} \\ y_0 e^{-2t} \end{bmatrix}$

Phase plane plot showing trajectories converging to the origin, with the y-axis converging faster than the x-axis.

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**"Phase Plane"**

**Example**

$$\frac{d}{dt} \begin{bmatrix} x \\ y \end{bmatrix} = \begin{bmatrix} -2 & 0 \\ 0 & -1 \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}$$

$$\begin{bmatrix} x \\ y \end{bmatrix}(0) = \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}$$

$$\dot{x} = -2x$$

$$\dot{y} = -y$$

$$\varphi \left( \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}, t \right) = \begin{bmatrix} x_0 e^{-2t} \\ y_0 e^{-t} \end{bmatrix}$$

$$x(0) = x_0$$

$$y(0) = y_0$$

$\begin{bmatrix} x_0 \\ y_0 \end{bmatrix} = \begin{bmatrix} 1 \\ 1 \end{bmatrix}$

$\begin{bmatrix} x_0 \\ y_0 \end{bmatrix} = \begin{bmatrix} 1 \\ 0 \end{bmatrix}$

$\begin{bmatrix} x_0 \\ y_0 \end{bmatrix} = \begin{bmatrix} 0 \\ -1 \end{bmatrix}$

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**"Phase Plane"**

**Example: "stable nodes"**

$$\frac{d}{dt} \begin{bmatrix} x \\ y \end{bmatrix} = \begin{bmatrix} a & 0 \\ 0 & b \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}$$

$$\begin{bmatrix} x \\ y \end{bmatrix}(0) = \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}$$

$$\dot{x} = ax$$

$$\dot{y} = by$$

$$\varphi \left( \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}, t \right) = \begin{bmatrix} x_0 e^{at} \\ y_0 e^{bt} \end{bmatrix}$$

$$x(0) = x_0$$

$$y(0) = y_0$$

$a < b < 0$

$a = b < 0$

$b < a < 0$

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**Math 5490**  
Dynamical Systems

**"Phase Plane"**

**Example: "unstable nodes"**

$$\frac{d}{dt} \begin{bmatrix} x \\ y \end{bmatrix} = \begin{bmatrix} a & 0 \\ 0 & b \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}$$

$$\begin{bmatrix} x \\ y \end{bmatrix}(0) = \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}$$

$$\dot{x} = ax$$

$$\dot{y} = by$$

$$\varphi \left( \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}, t \right) = \begin{bmatrix} x_0 e^{at} \\ y_0 e^{bt} \end{bmatrix}$$

$$x(0) = x_0$$

$$y(0) = y_0$$

$\alpha > \beta > 0$

$\alpha = \beta > 0$

$0 < \alpha < \beta$

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**Math 5490**  
Dynamical Systems

**"Phase Plane"**

**Example: "saddles"**

$$\frac{d}{dt} \begin{bmatrix} x \\ y \end{bmatrix} = \begin{bmatrix} a & 0 \\ 0 & b \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}$$

$$\begin{bmatrix} x \\ y \end{bmatrix}(0) = \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}$$

$$\dot{x} = ax$$

$$\dot{y} = by$$

$$\varphi \left( \begin{bmatrix} x_0 \\ y_0 \end{bmatrix}, t \right) = \begin{bmatrix} x_0 e^{at} \\ y_0 e^{bt} \end{bmatrix}$$

$$x(0) = x_0$$

$$y(0) = y_0$$

unstable manifold

stable manifold

$\alpha < 0 < \beta$

$\beta < 0 < \alpha$

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